

Improving Understanding and Prediction of High Impact Weather Associated with Low-Topped Severe Convection in the Southeastern U.S.

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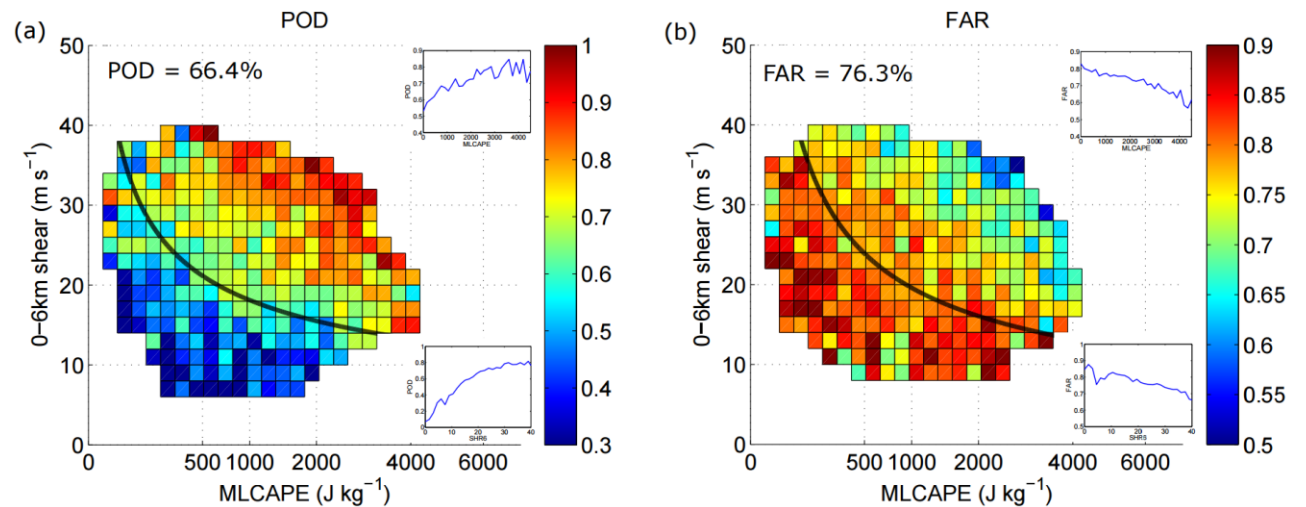
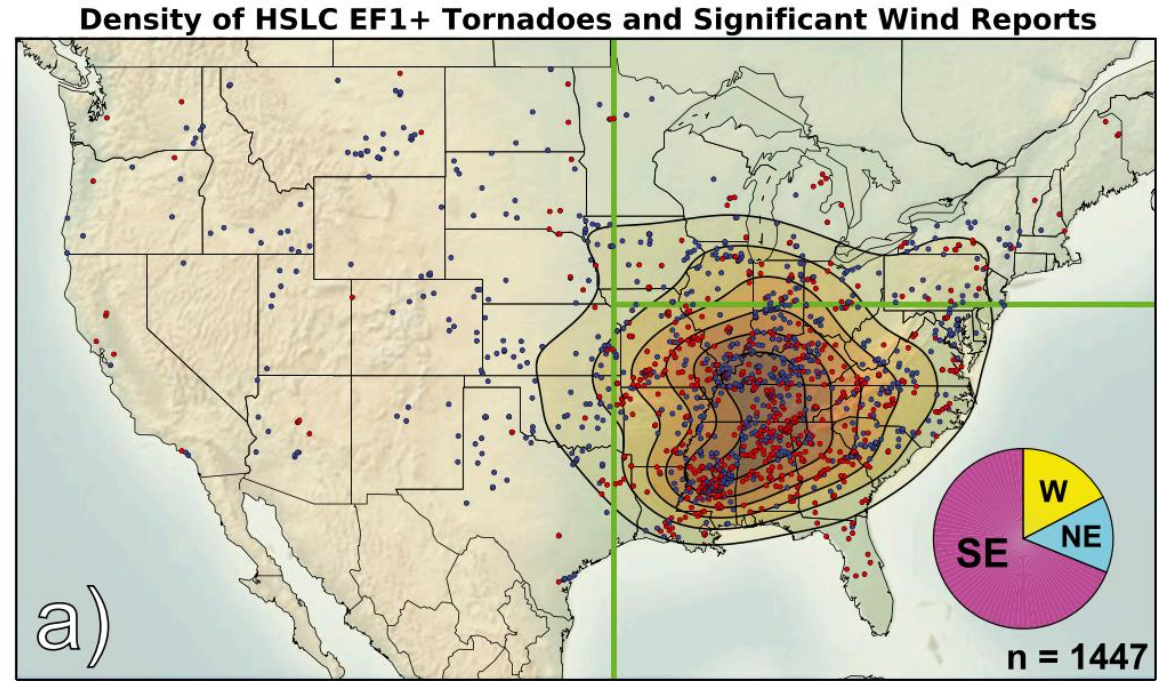
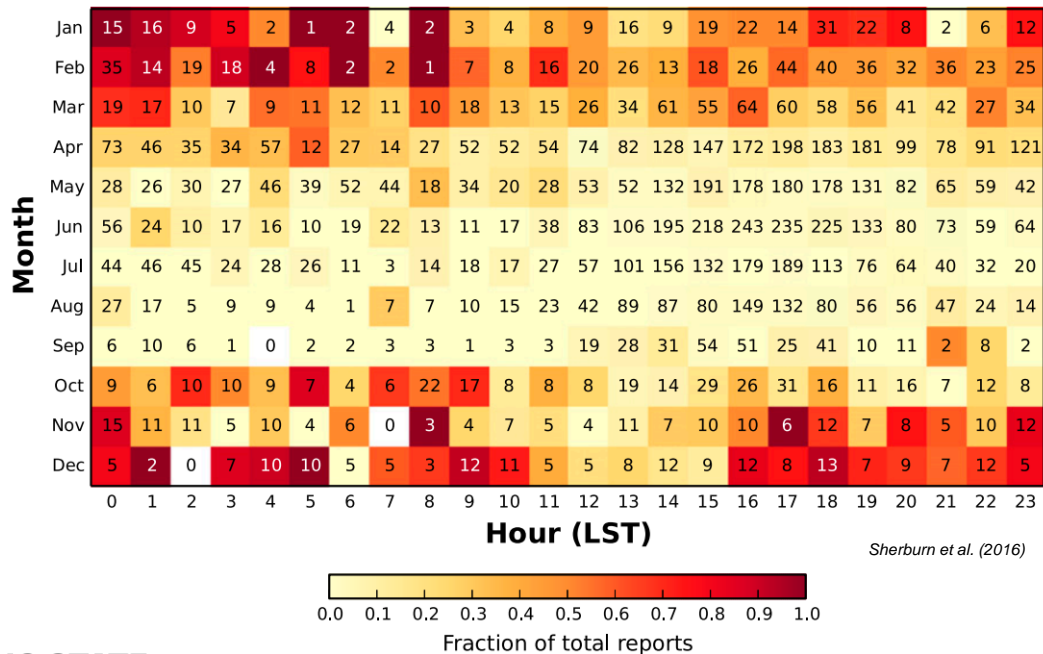
Collaborators: ESRL; EMC; SPC; WFOs Birmingham, Blacksburg, Charleston (SC), Columbia, Greer, Huntsville, Morehead City/Newport, Peachtree City, Raleigh, Sterling, Tallahassee, Wakefield, Wilmington (NC), Wilmington (OH)

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Problem overview

High-shear, low-CAPE (HSLC*) severe convection is a considerable forecasting challenge across the eastern U.S., particularly during the cool season and overnight



*Here, defined as $SBCAPE \leq 500 \text{ J kg}^{-1}$, $MUCAPE \leq 1000 \text{ J kg}^{-1}$, and $0\text{-}6 \text{ km bulk wind difference} \geq 18 \text{ m s}^{-1}$

Project overview

Five components:

- Composite maps and parameters:
 - Determine typical features associated with severe/nonsevere HSLC events (Sherburn et al. 2017)
 - Assess operational utility of existing and new forecasting parameters (Sherburn et al. 2017)
- Process studies 1: Case simulations to study mesoscale/synoptic scale evolution (King et al. 2017)
- NWP studies: Case simulations to investigate resolution requirements
- Process studies 2: Idealized simulations to study convective-scale dynamics
- Statistical studies: Dynamical-statistical downscaling to investigate predictability

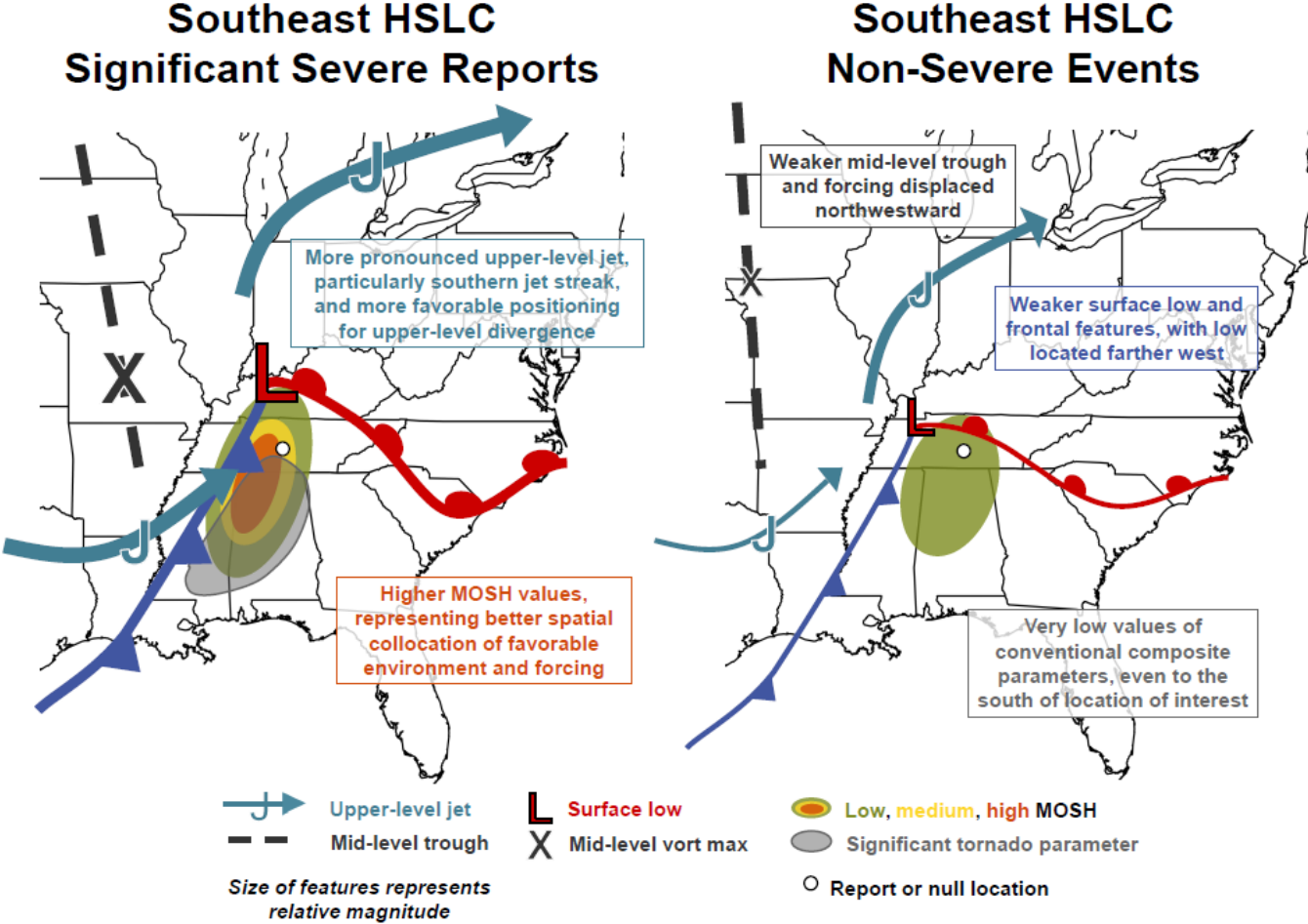
HSLC composites: Key points

Created using NARR data and severe reports versus false alarm warnings

Stronger, more closely collocated features are conducive to severe HSLC events

Release of potential instability and/or strong low-level θ_e advection are responsible for rapid destabilization immediately ahead of HSLC convection

Low-level lapse rates and shear vector magnitudes remain skillful



HSLC composites: Updated forecasting parameter

$$\text{MOSH} = \frac{(\text{LLLR} - 4 \text{ K km}^{-1})^2}{4 \text{ K}^2 \text{ km}^{-2}} \times \frac{(\text{S15MG} - 8 \text{ m s}^{-1})}{10 \text{ m s}^{-1}} \times \frac{(\text{MAXTEVV} + 10 \text{ K Pa km}^{-1} \text{ s}^{-1})}{9 \text{ K Pa km}^{-1} \text{ s}^{-1}}$$

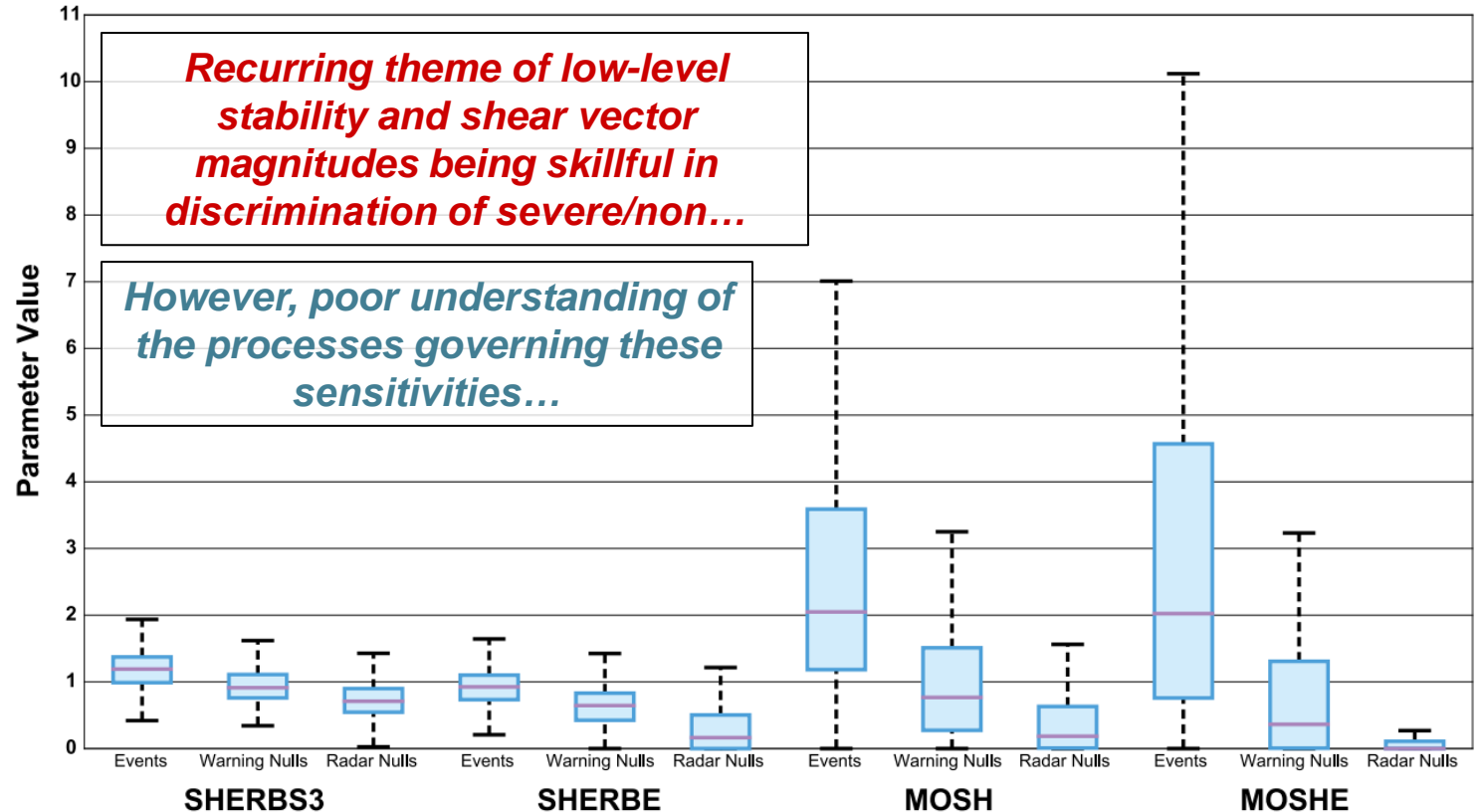
$$\text{MOSHE} = \frac{(\text{LLLR} - 4 \text{ K km}^{-1})^2}{4 \text{ K}^2 \text{ km}^{-2}} \times \frac{(\text{S15MG} - 8 \text{ m s}^{-1})}{10 \text{ m s}^{-1}} \times \frac{(\text{ESHR} - 8 \text{ m s}^{-1})}{10 \text{ m s}^{-1}} \times \frac{(\text{MAXTEVV} + 10 \text{ K Pa km}^{-1} \text{ s}^{-1})}{9 \text{ K Pa km}^{-1} \text{ s}^{-1}}$$

LLLR: 0-3 km lapse rate

S15MG: 0-1.5 km shear vector magnitude

ESHR: Effective shear magnitude

MAXTEVV: Maximum $d\theta_e/dz * \omega$ product from 0-2 km through 0-6 km, calculated at 0.5 km intervals (positive: unstable/upward motion)



HSLC composites: MOSHE on SPC Mesoanalysis

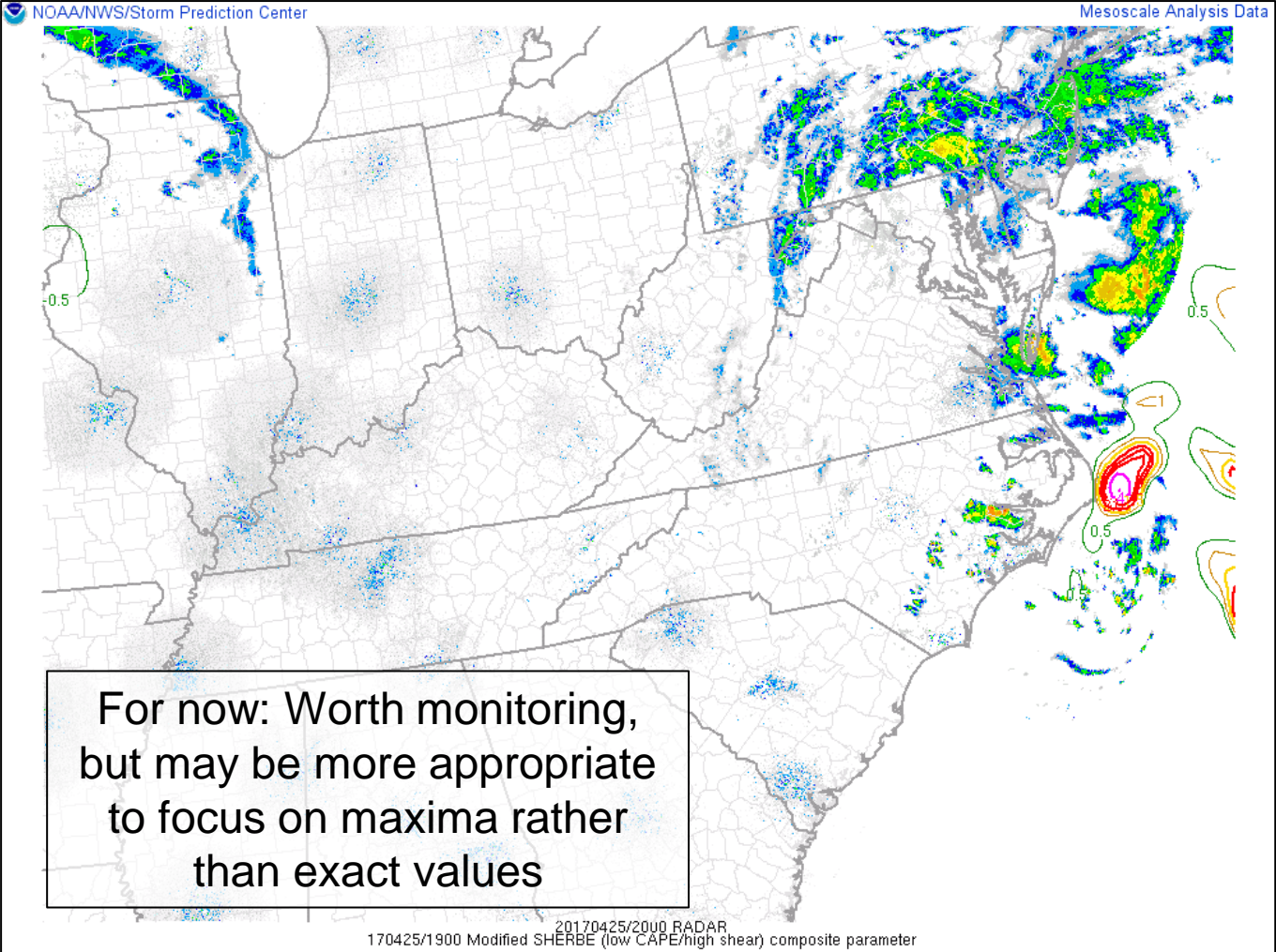
$$\text{MOSH} = \frac{(\text{LLLR} - 4 \text{ K km}^{-1})^2}{4 \text{ K}^2 \text{ km}^{-2}} \times \frac{(\text{S15MG} - 8 \text{ m s}^{-1})}{10 \text{ m s}^{-1}} \times \frac{(\text{MAXTEVV} + 10 \text{ K Pa km}^{-1} \text{ s}^{-1})}{9 \text{ K Pa km}^{-1} \text{ s}^{-1}}$$

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← Adjustments potentially necessary; NARR values biased low

Could “cap” terms to limit contributions or reevaluate normalization values

- LLLR:** 0-3 km lapse rate
- S15MG:** 0-1.5 km shear vector magnitude
- ESHR:** Effective shear magnitude
- MAXTEVV:** Maximum $d\theta_e/dz * \omega$ product from 0-2 km through 0-6 km, calculated at 0.5 km intervals (positive: unstable/upward motion)



Process studies: Case study selection

Requirements:

- At least “slight” risk for severe convection
 - SPC mesoanalysis CAPE $\leq 1000 \text{ J kg}^{-1}$
 - 0-3 km shear $\geq 18 \text{ m s}^{-1}$

6 non-severe events (no storm reports)

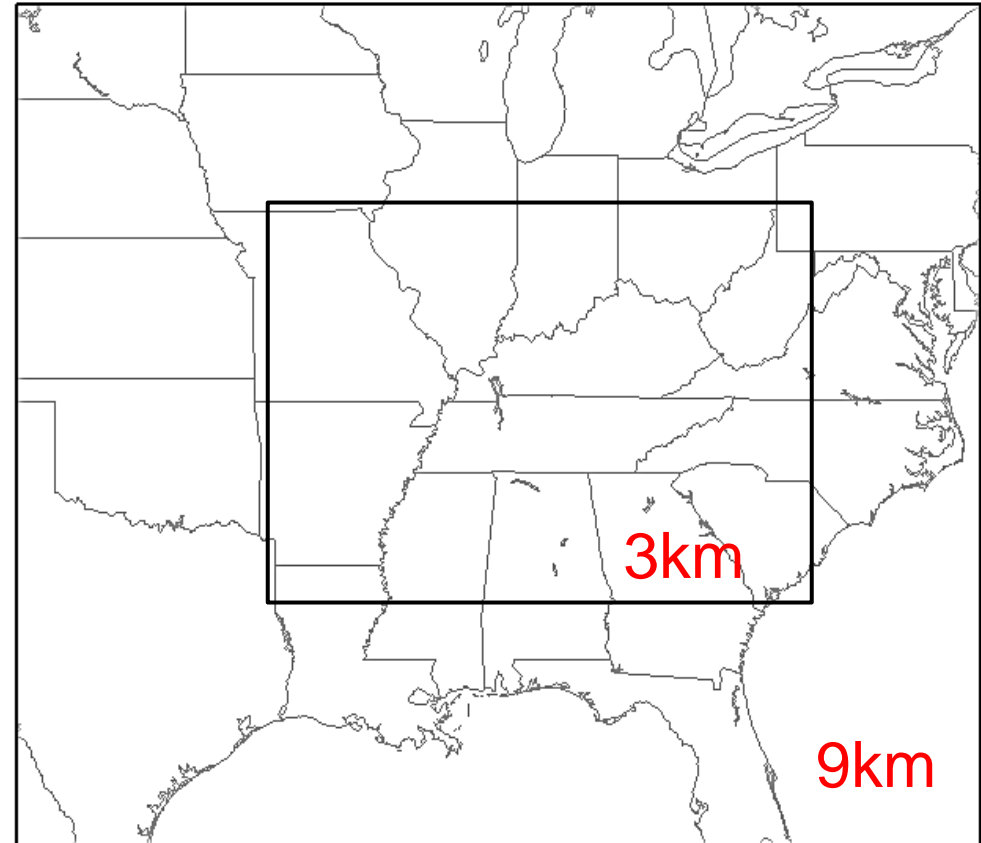
11 severe events (multiple reports)

Simulations run with ARW-WRF, v3.5.1

50 vertical levels

6-h NAM 12-km analyses as IC/LBC

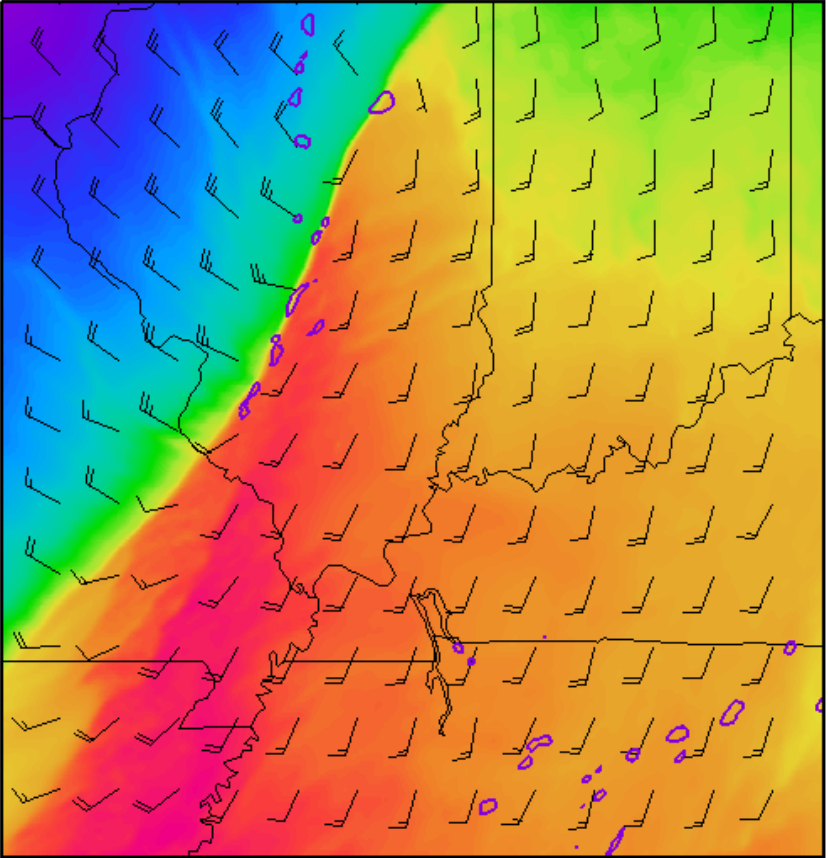
At least 30 hrs simulation time



Process studies: Simulated environments

Jan. 29, 2008: 2pm – 6pm

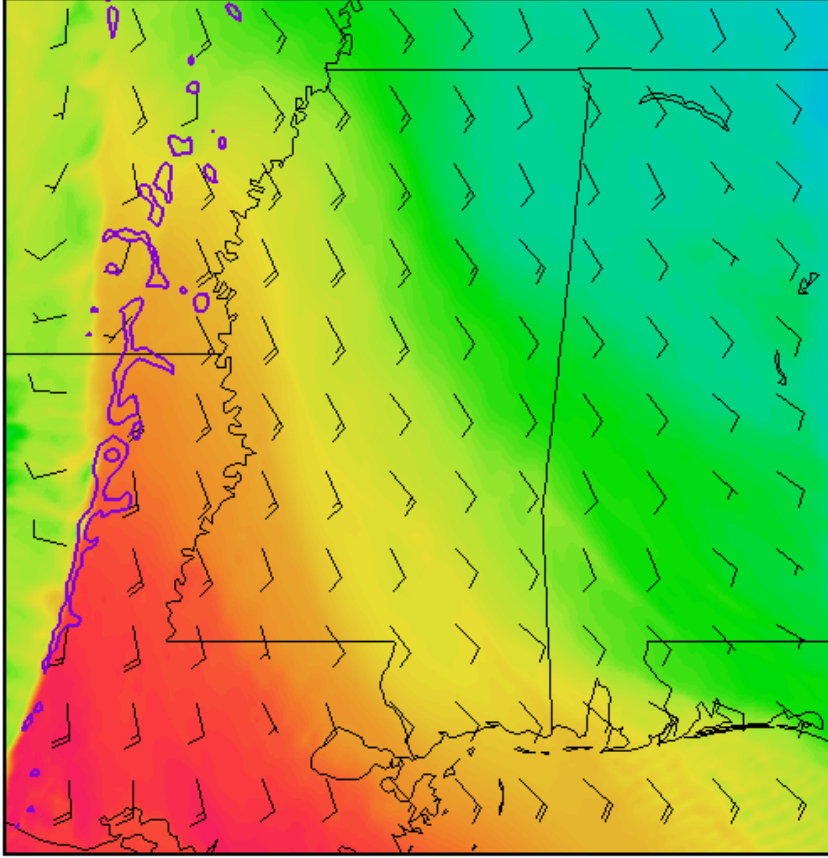
Surface θ_e , 10 m wind barbs [kts], 40 dBZ contour



Severe

Dec. 22, 2007: 4pm – 8pm

Surface θ_e , 10 m wind barbs [kts], 40 dBZ contour



Nonsevere

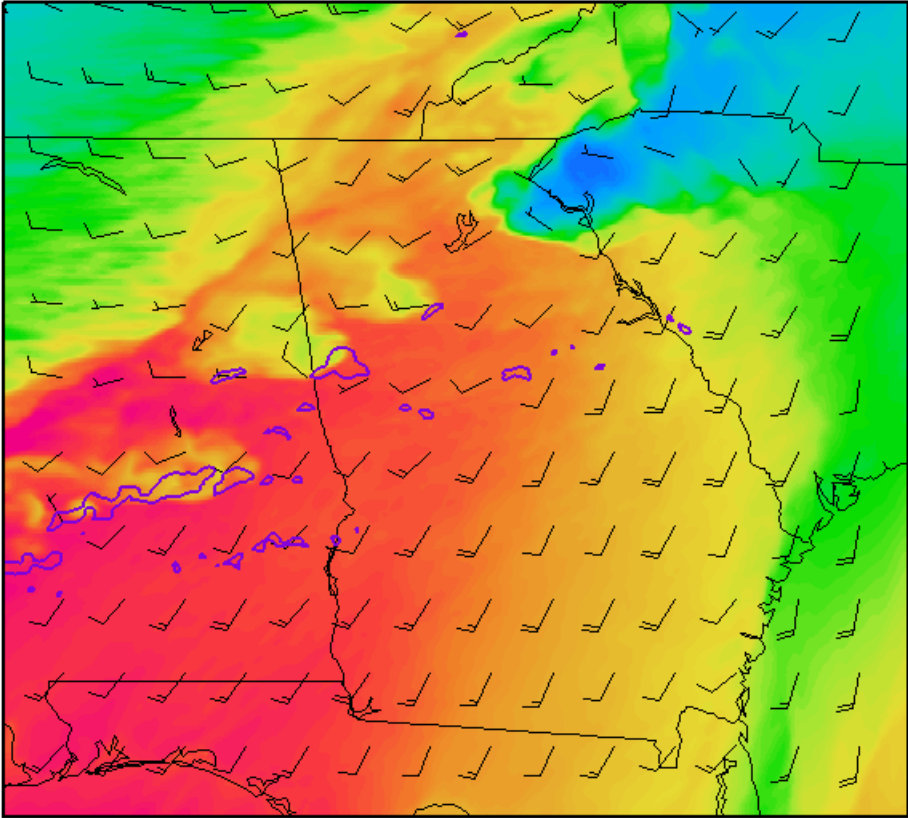


Linear cold frontal boundary

Process studies: Simulated environments

Feb. 18, 2009: 5pm – 9pm

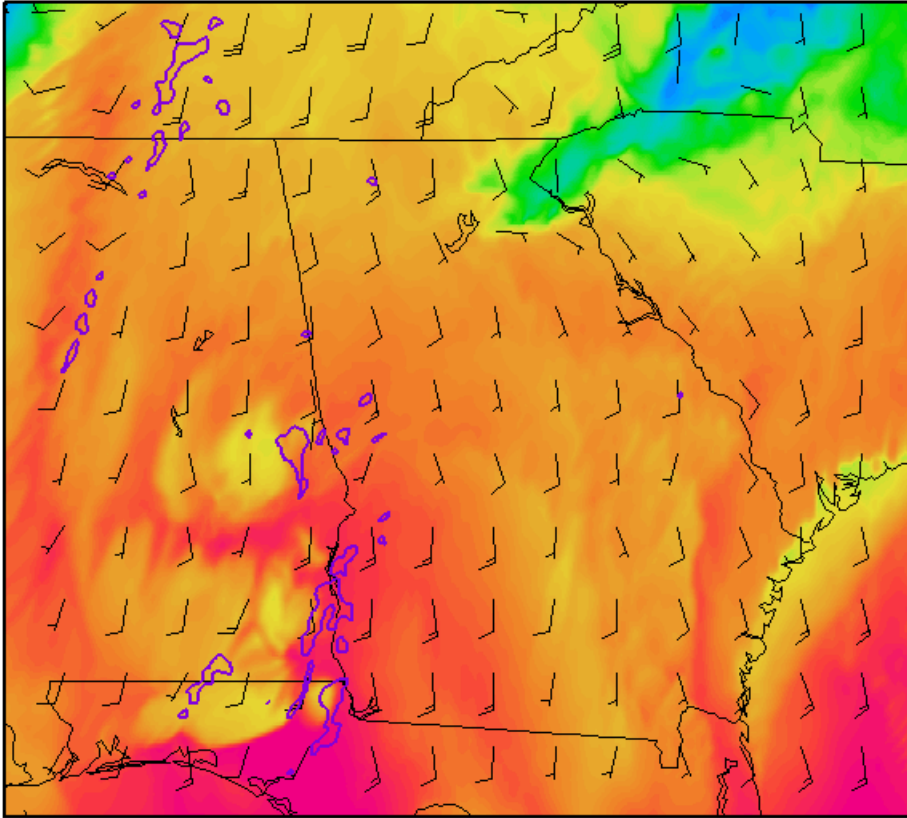
Surface θ_e , 10 m wind barbs [kts], 40 dBZ contour



Severe

Jan 26, 2012: 5pm – 9pm

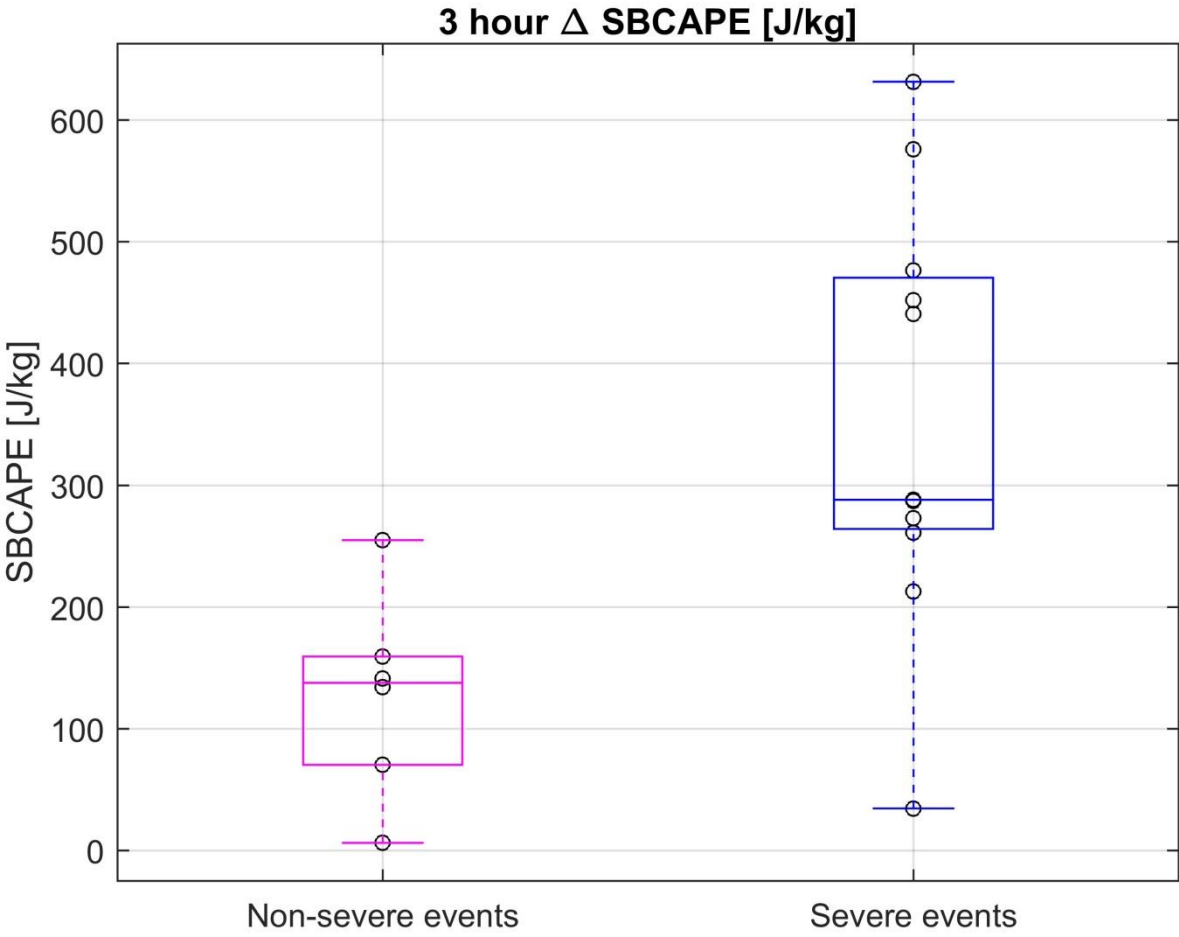
Surface θ_e , 10 m wind barbs [kts], 40 dBZ contour



Nonsevere

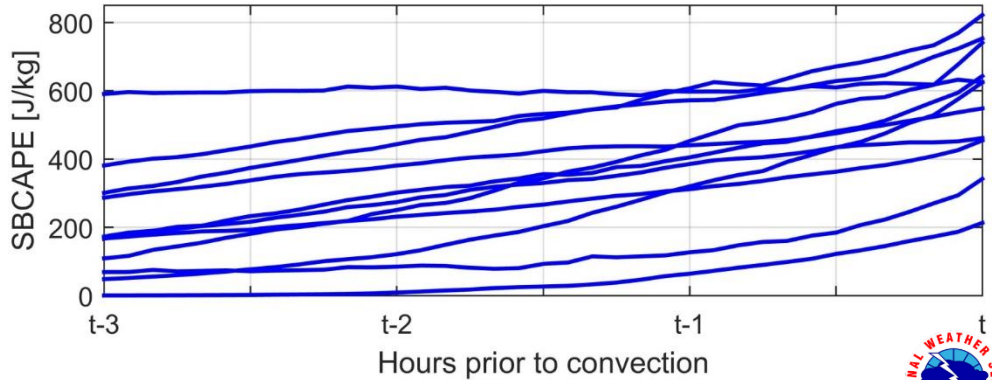
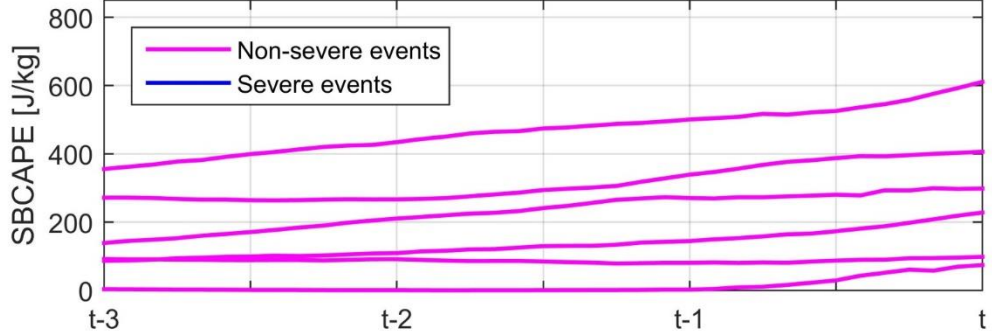
Isolated convection

Process studies: CAPE increases



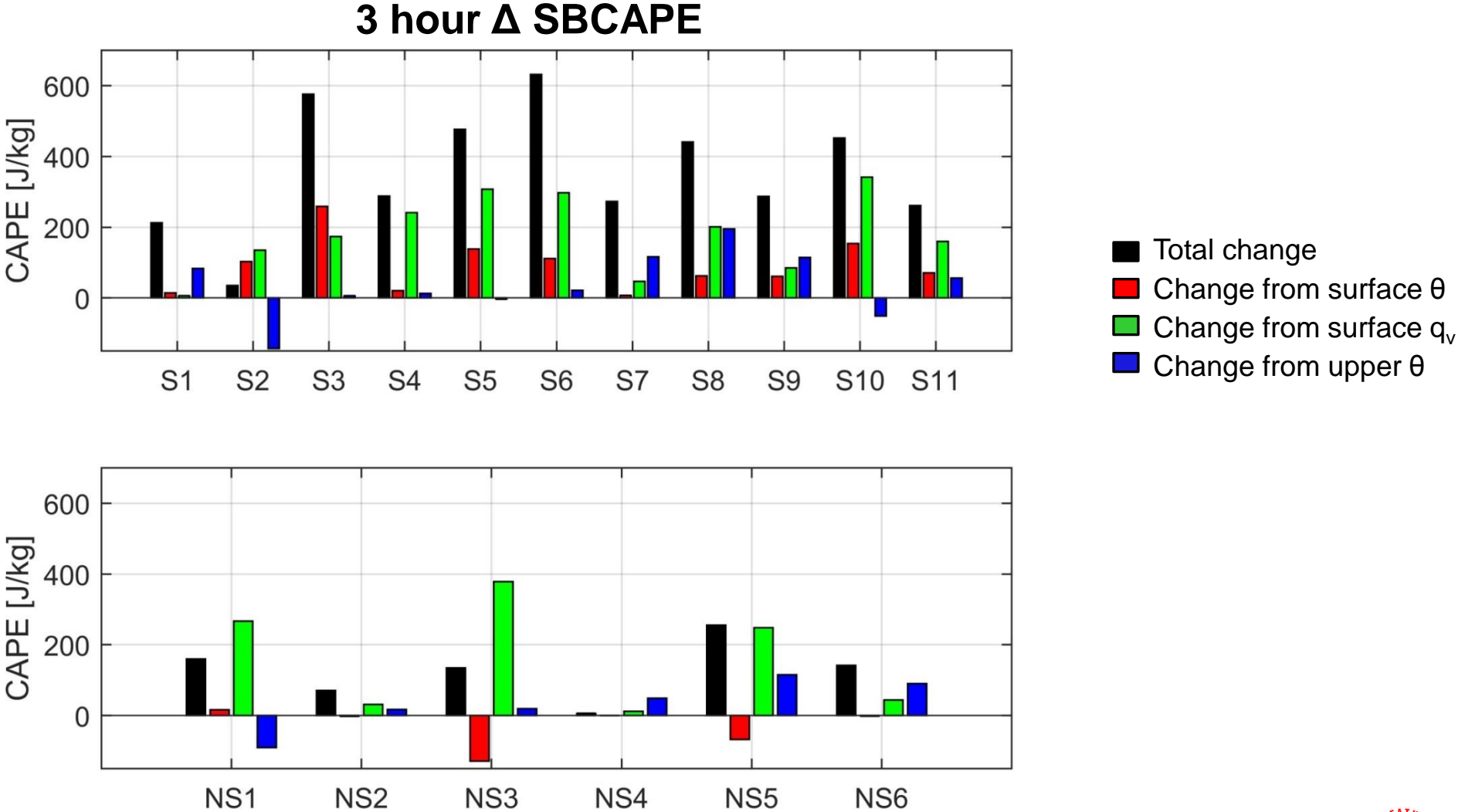
CAPE increases could arise from:

- Increased surface temperature
- Increased surface moisture
- Decreased temperature aloft

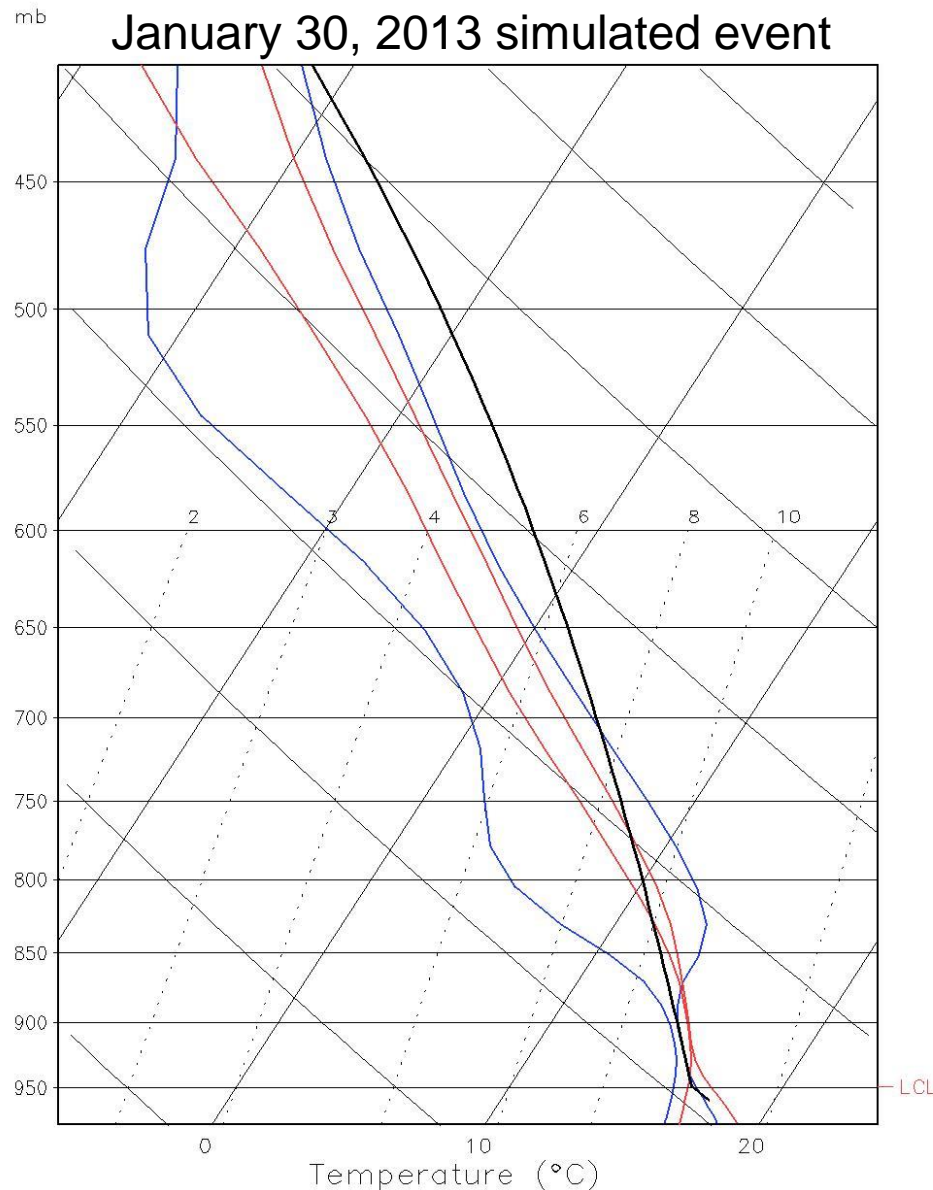


Goal: Determine which processes are most important by calculating contributions to CAPE from each process

Process studies: CAPE increases



Process studies: Potential instability



Potentially unstable sounding (θ_e decreasing with height)

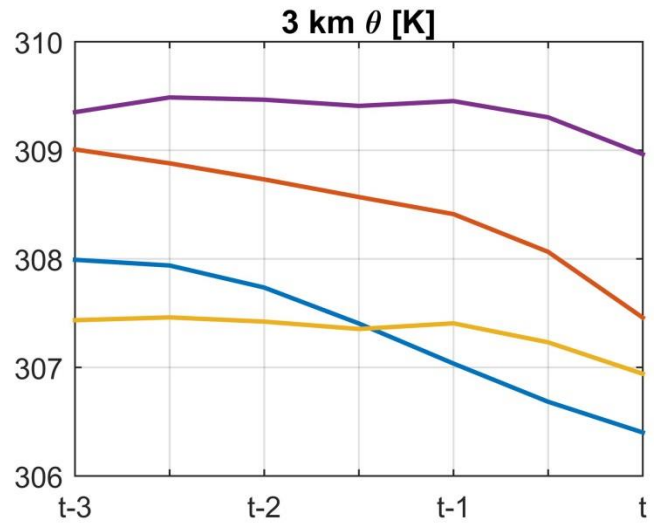
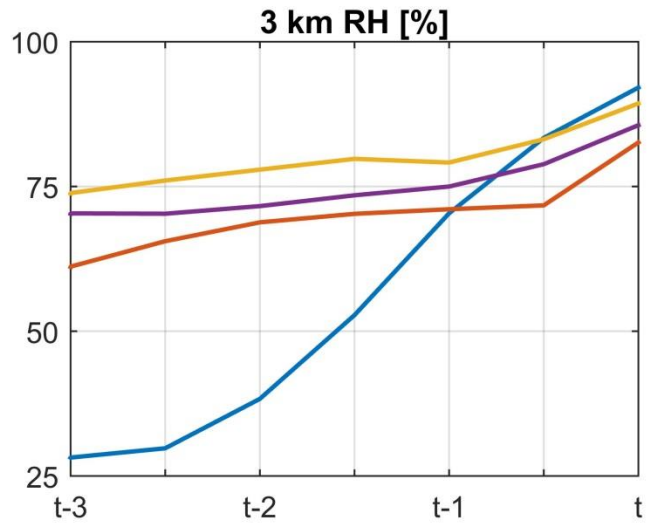
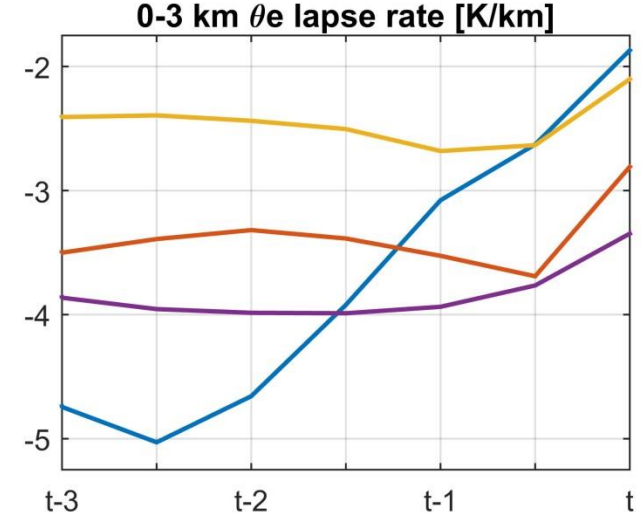
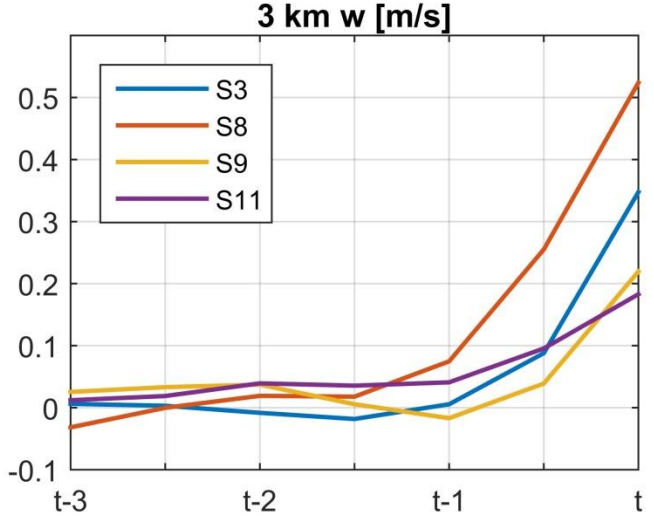
Lifting = cooling and moistening

Potential instability released, increase in CAPE

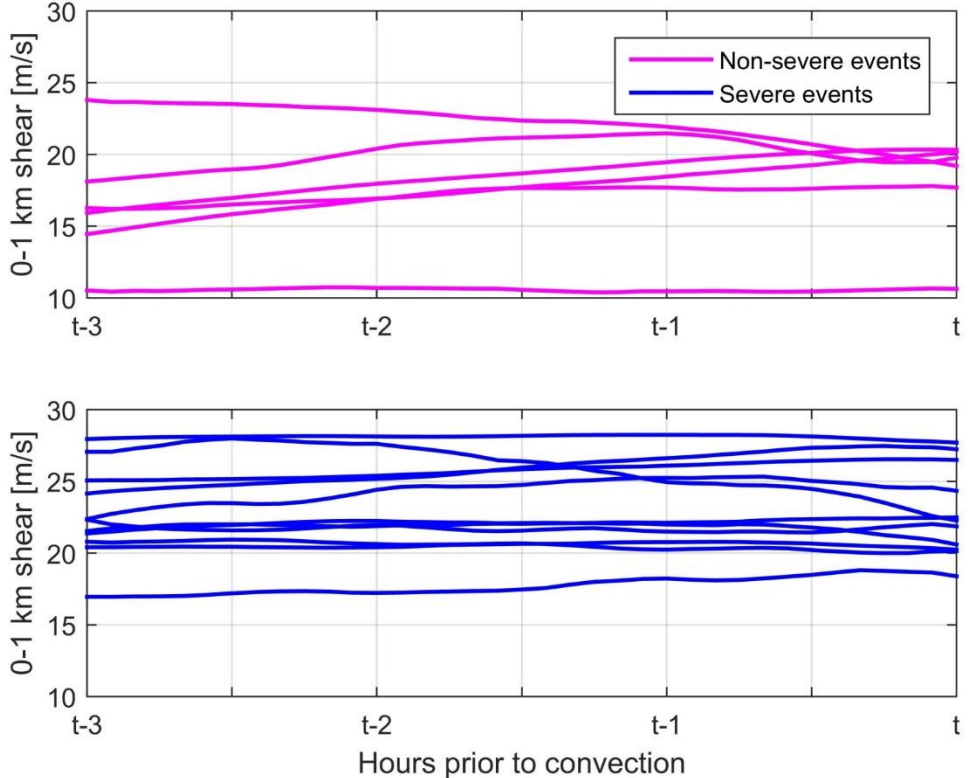
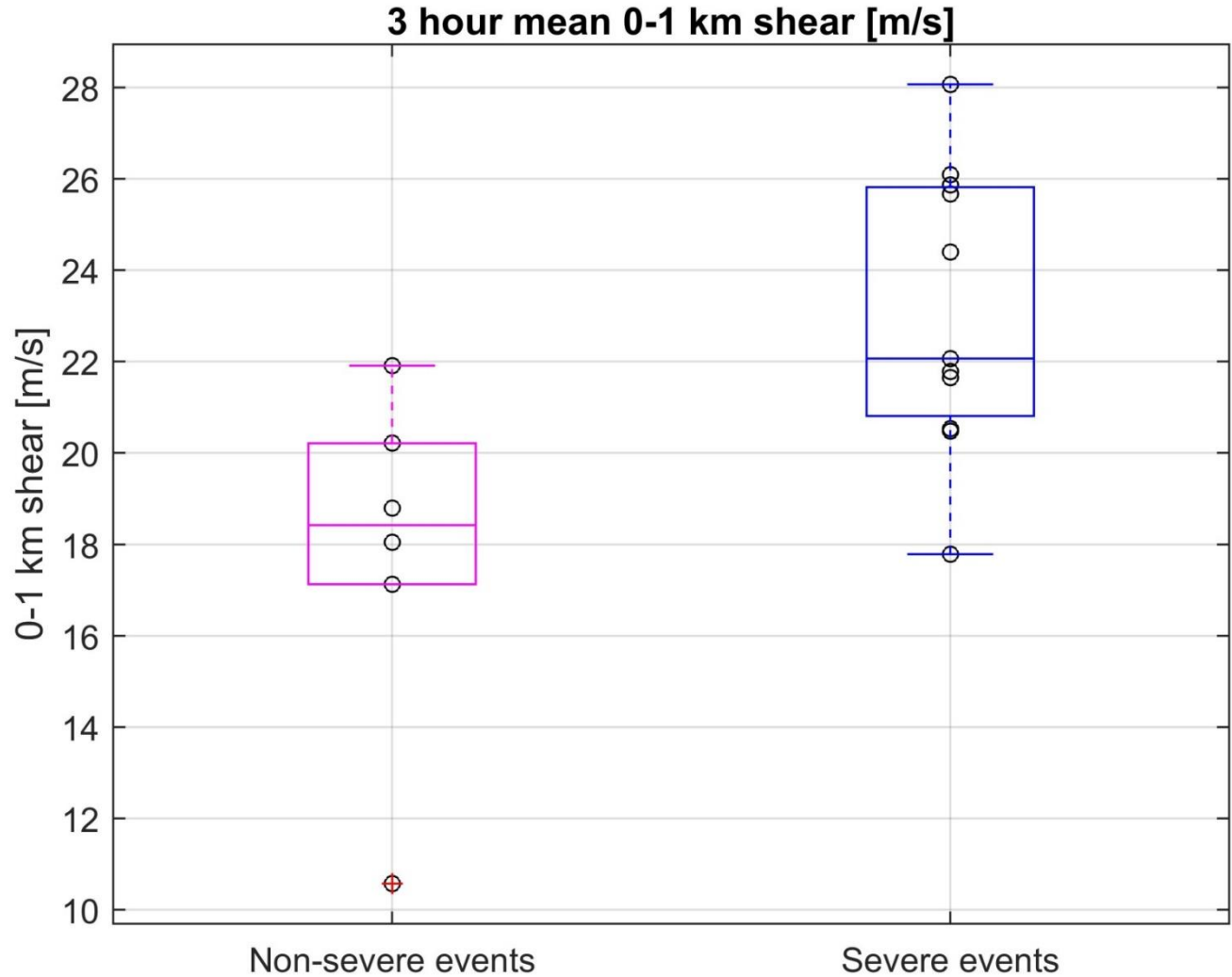
3 hours prior to convection

Just prior to convection (parcel path)

Process studies: Synoptic ascent

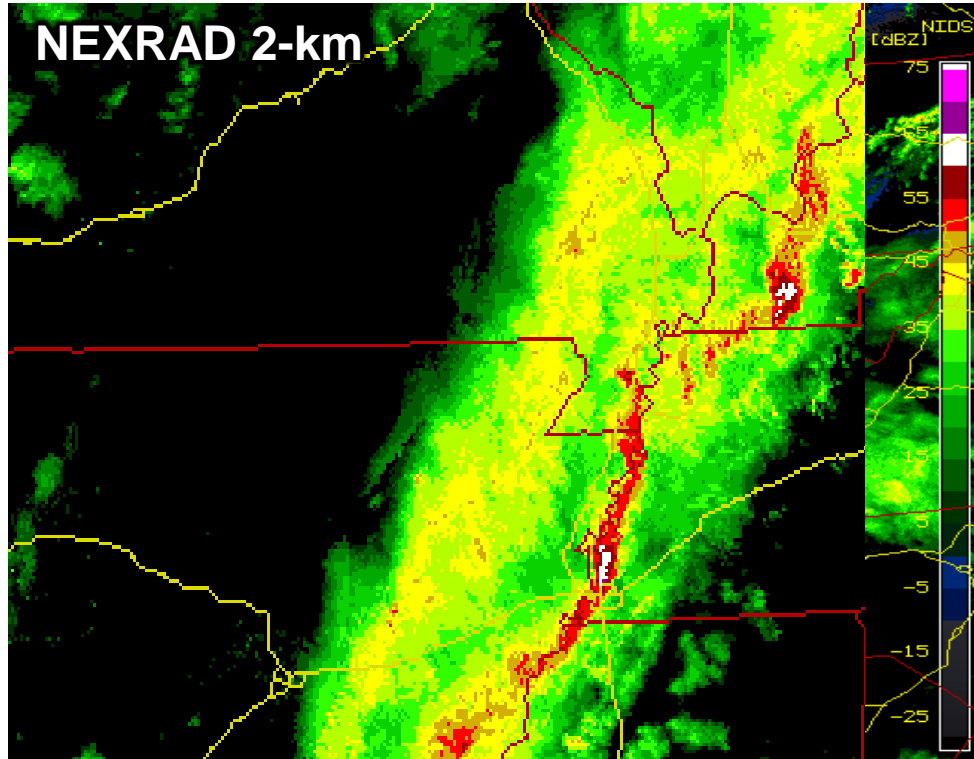


Process studies: Low-level shear

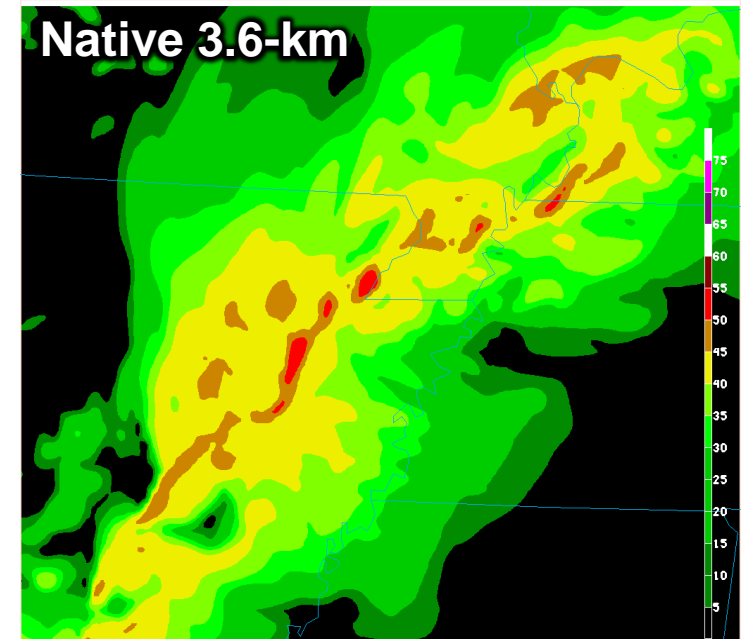


Discrimination between severe/nonsevere...
...but relatively constant over time

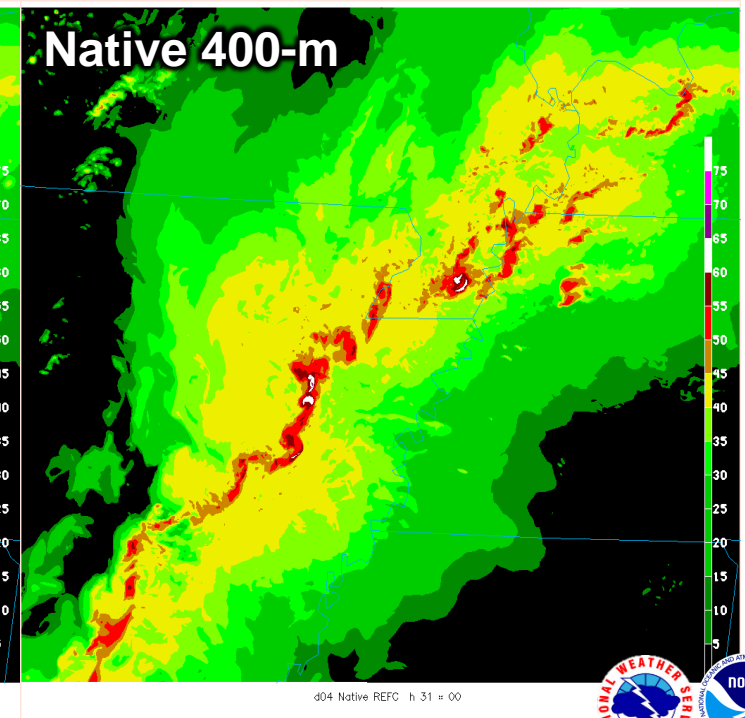
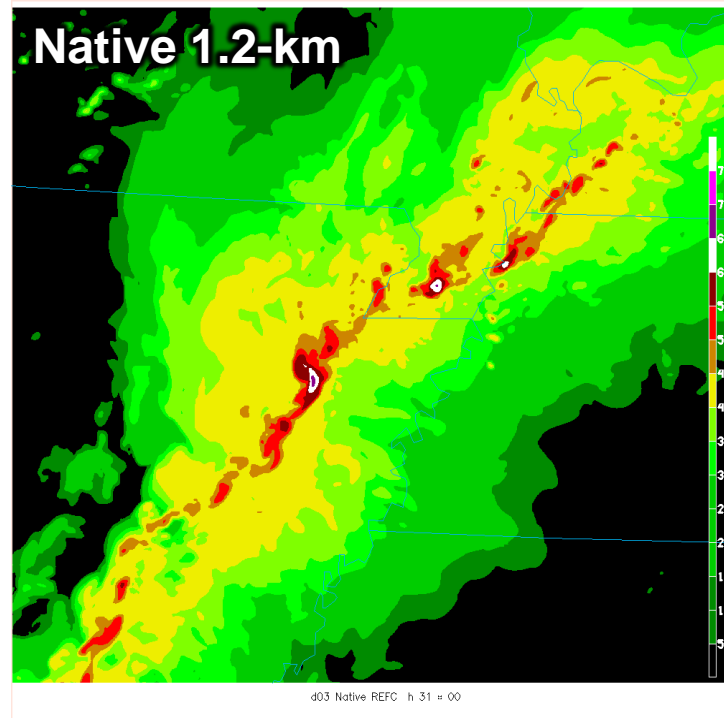
NWP studies: Resolution differences



Convective characteristics well-handled at all resolutions, though finer resolution obviously captures more detail



Differences clearer when evaluating fields such as 10-m wind speeds...



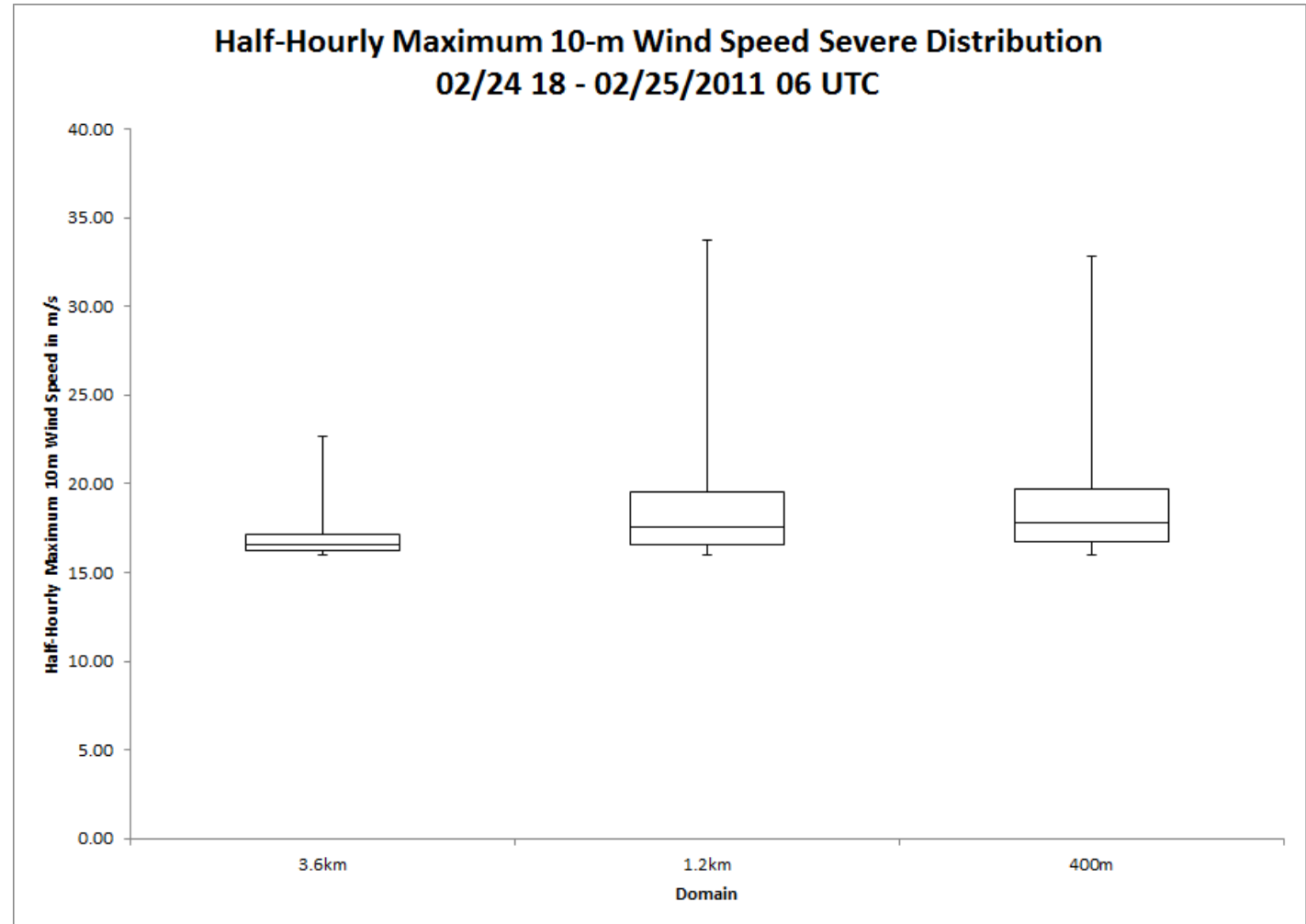
NWP studies: Resolution differences

Fairly large difference between 3.6-km and 1.2-km grid spacing, particularly when compared to jump from 1.2-km to 400-m

Similar findings for second case investigated

Suggests convergence of solution between 3.6-km and 1.2-km grid spacing, at least for sensible hazards

Work continues; more cases needed to corroborate initial findings

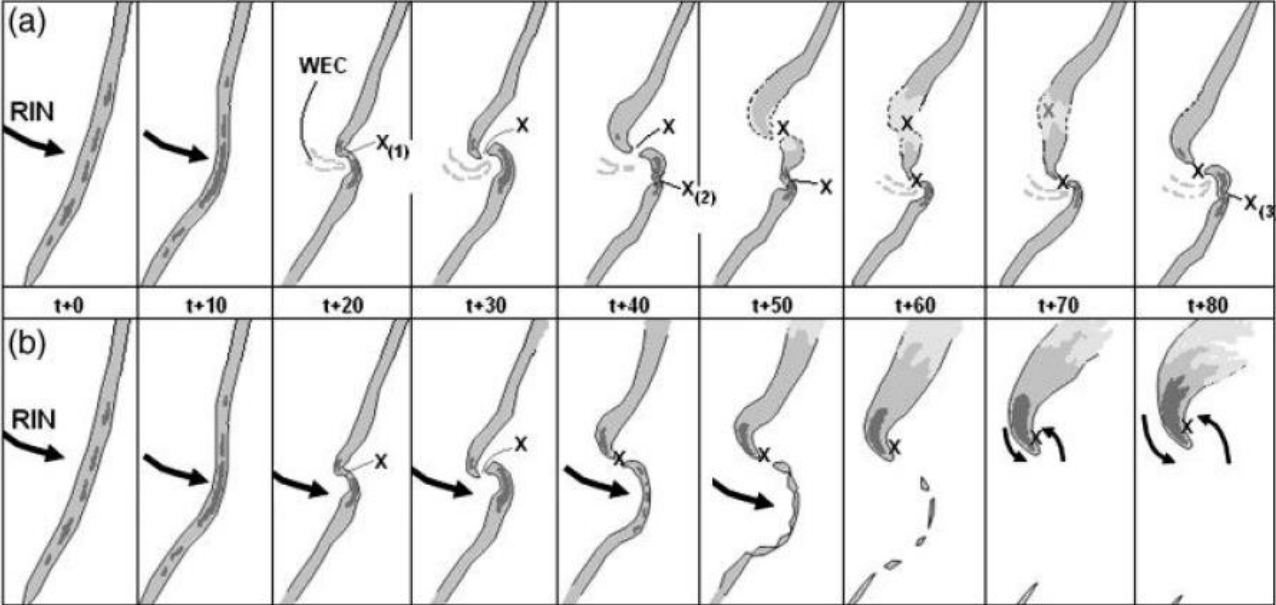
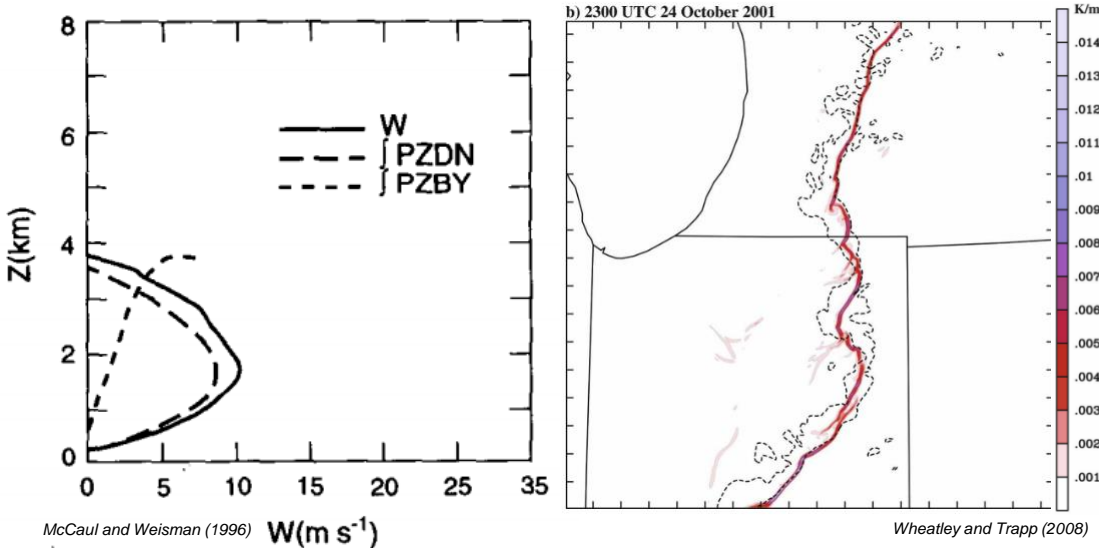


Idealized simulations: Motivation

Poor radar resolution and discrimination

Some potential radar precursors (such as broken-S, right), but high associated FAR

Rapid destabilization



Clark (2011)

Prior HSLC simulations mainly tropical mini-supercells with insufficient resolution

Very few HSLC QLCS studies, limited in scope

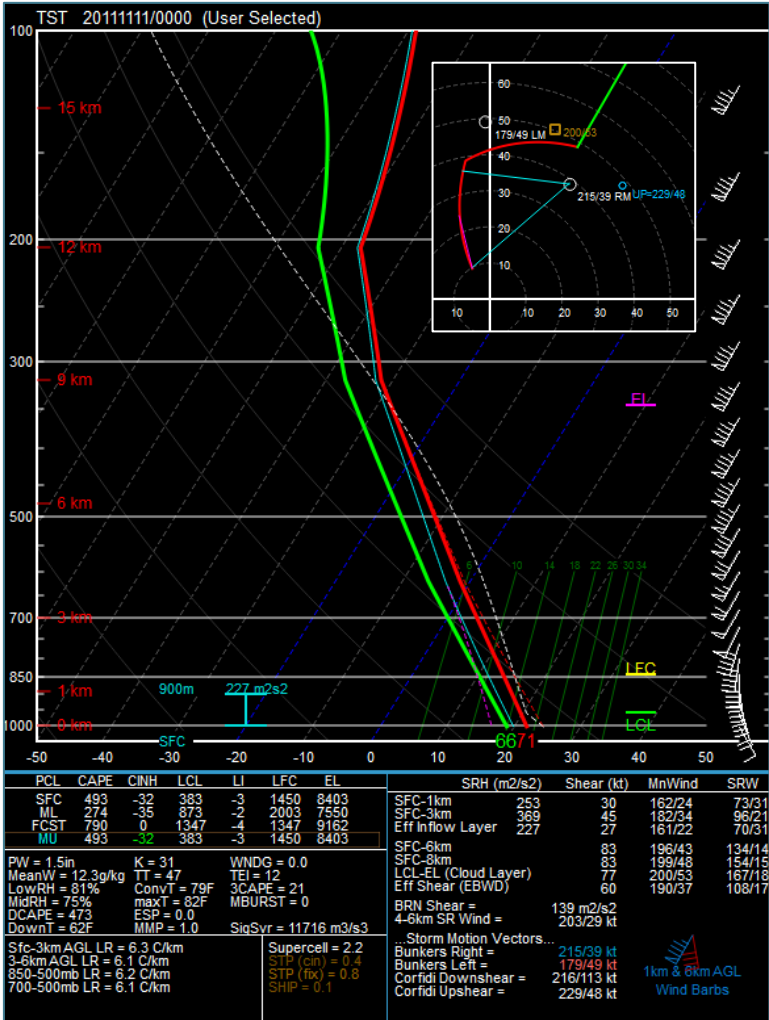
QLCS mesovortex genesis mechanisms uncertain

Idealized simulations: Overview

Two primary goals:

1. Determine environmental parameter space within HSLC environments where long-lived, strong, low-level vortices capable of producing severe hazards are likely
2. Determine precursors for the development of strong, low-level vortices to determine the dynamics governing documented sensitivities

We must first understand the links in the chain that extend from the development of HSLC convection to the development of strong, near-surface vortices therein before we can assess which links are broken or missing.

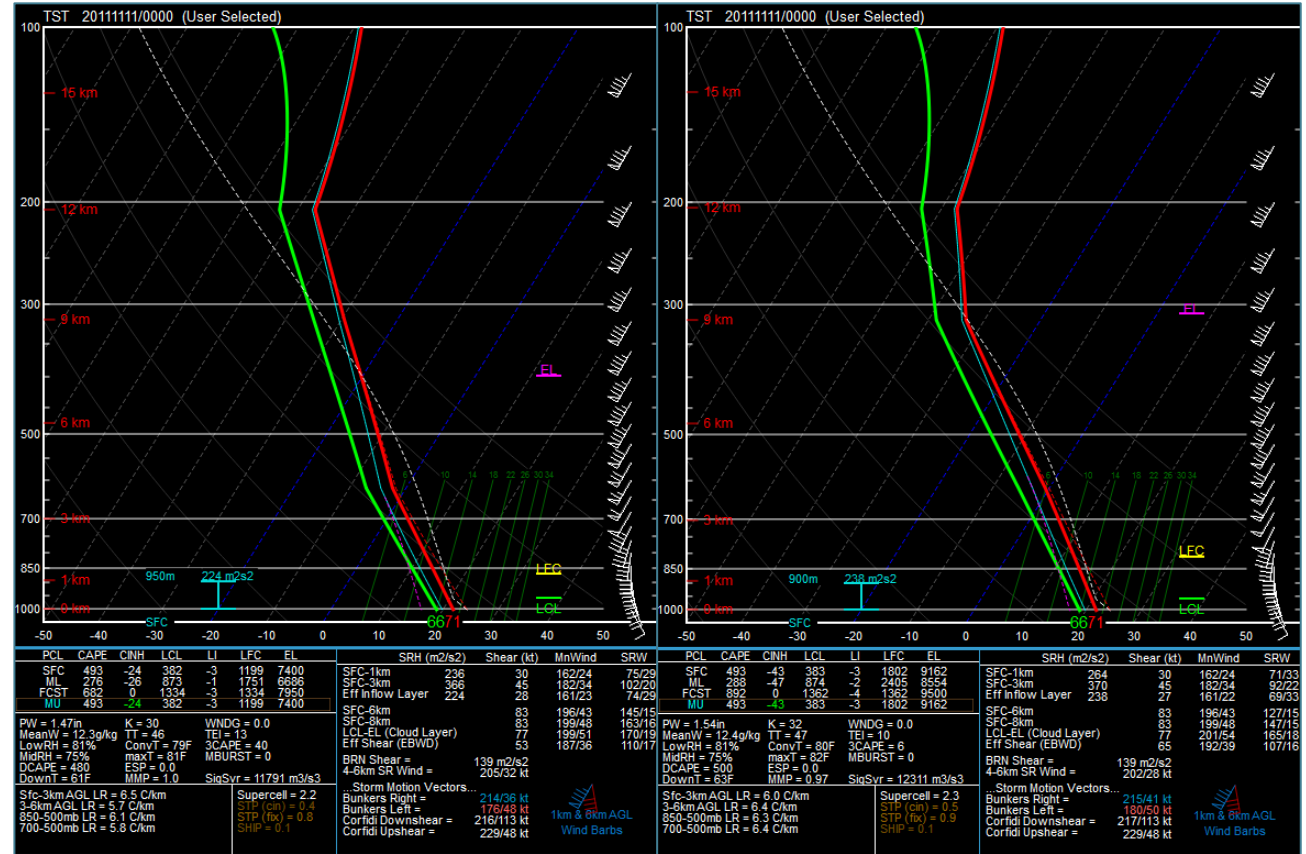
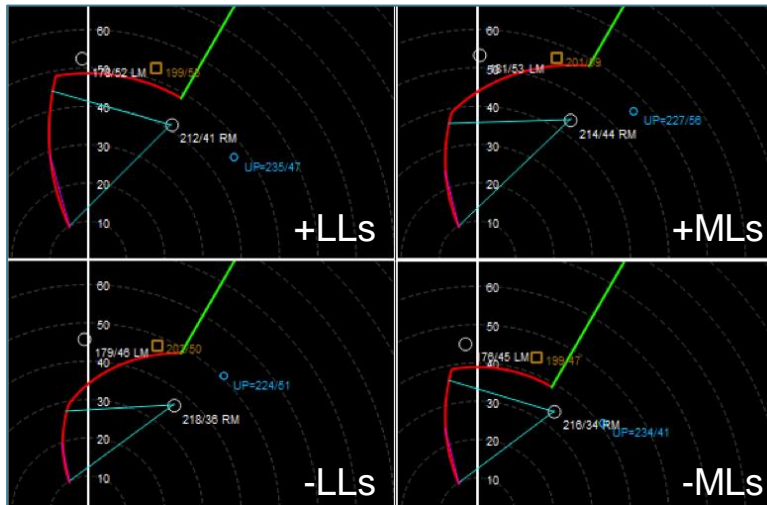


Control base-state environment

Idealized simulations: Sensitivity tests

Based on prior environmental studies and skill tests leading to the development of the SHERB/MOSH

Focus on varying low-level CAPE (here, equivalent to low-level lapse rates) and low/mid-level shear vector magnitudes



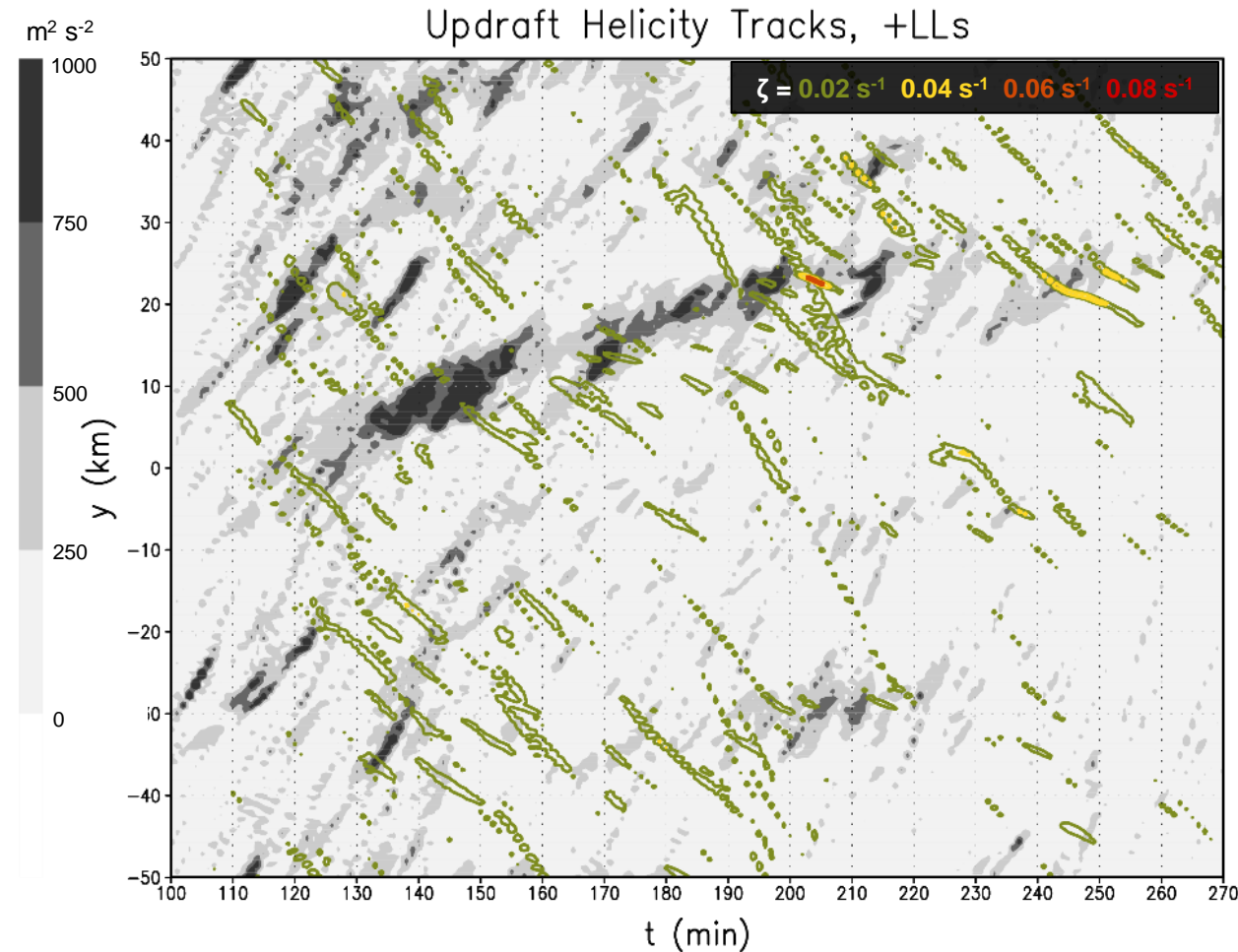
Increased low-level CAPE (+LLc)

Decreased low-level CAPE (-LLc)

Idealized simulations: Informed hypotheses

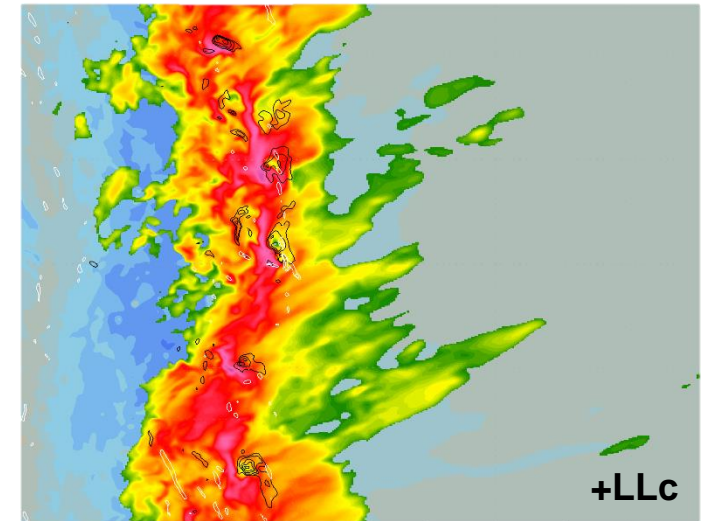
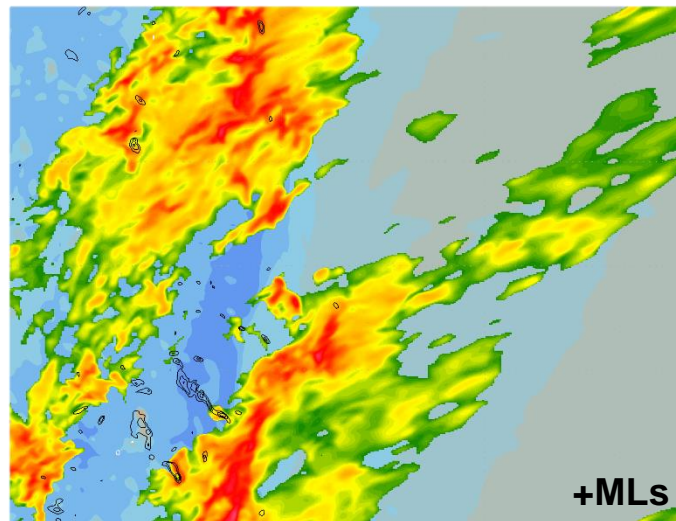
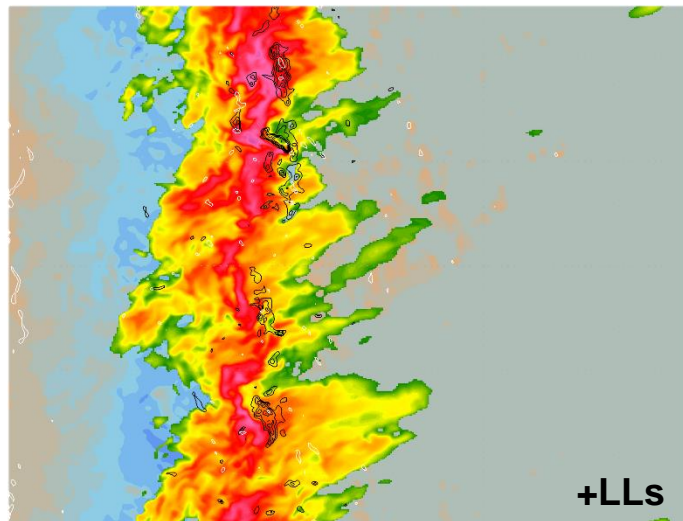
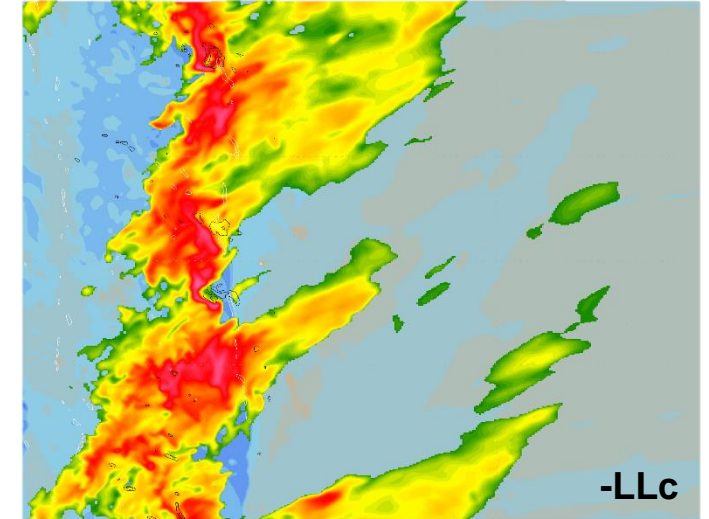
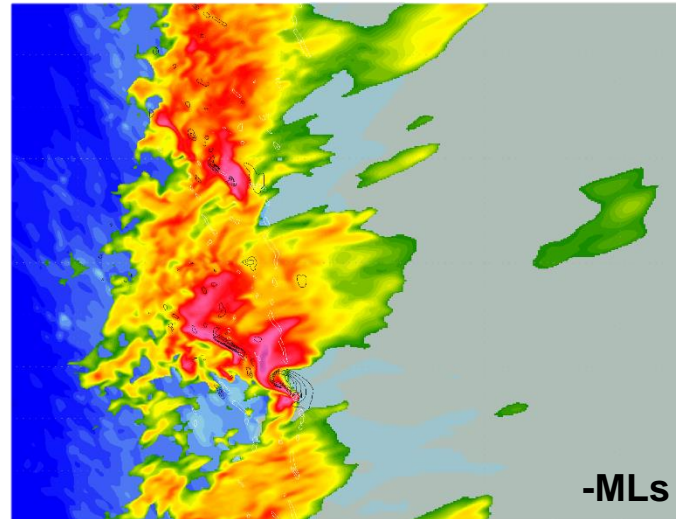
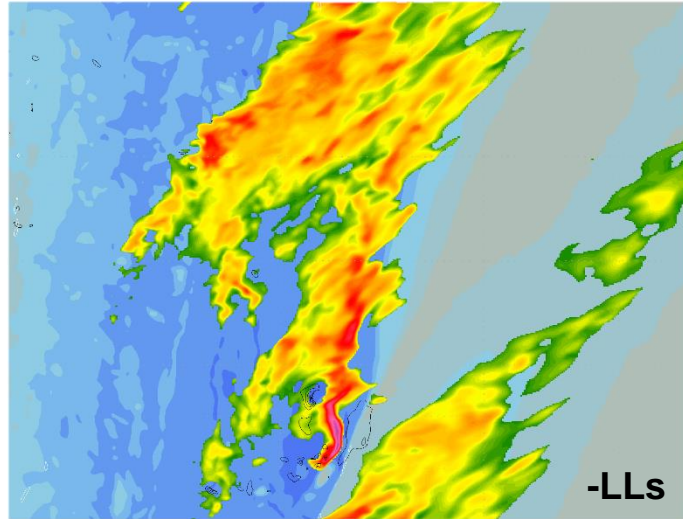
A strong, low-level updraft ($\sim 20 \text{ m s}^{-1}$ in the lowest 2 km) is a necessary but insufficient precursor to low-level vortexgenesis

Increasing this variable...	...will lead to these results
Low-level shear	<ul style="list-style-type: none"> Increased <i>number</i> of strong low-level updrafts and near-surface ζ centers Increased probability of producing a strong, near-surface vortex
Low-level CAPE	<ul style="list-style-type: none"> Increased <i>number and magnitude</i> of strong low-level updrafts Increased probability of producing a strong, near-surface vortex



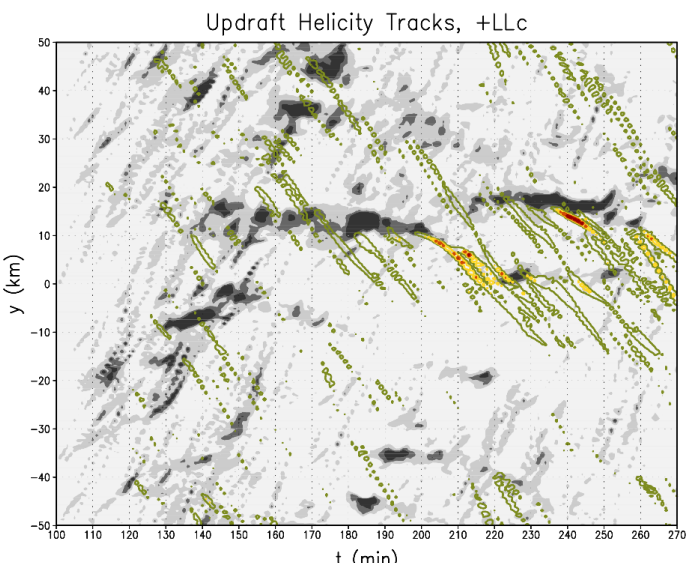
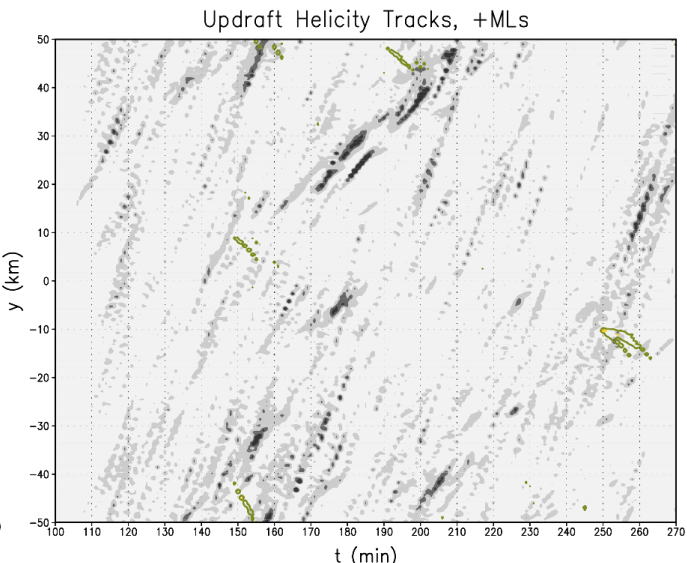
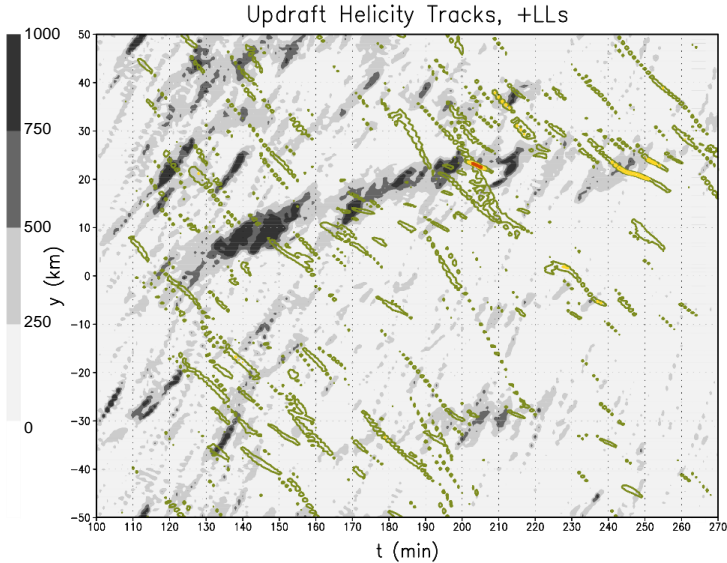
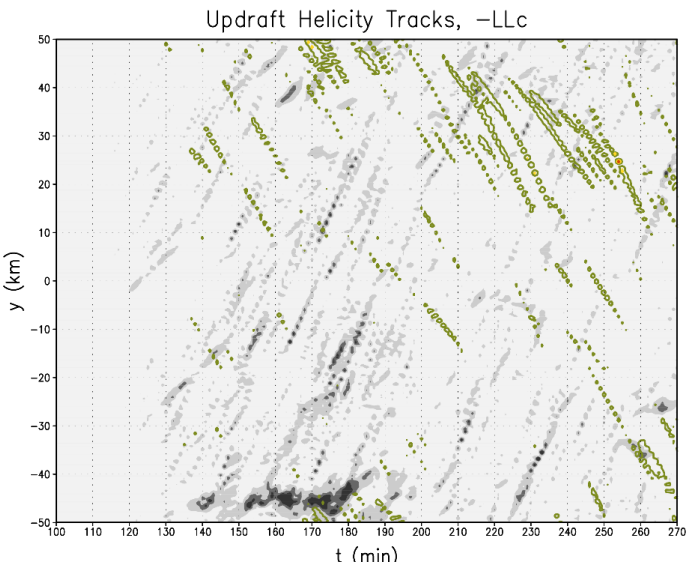
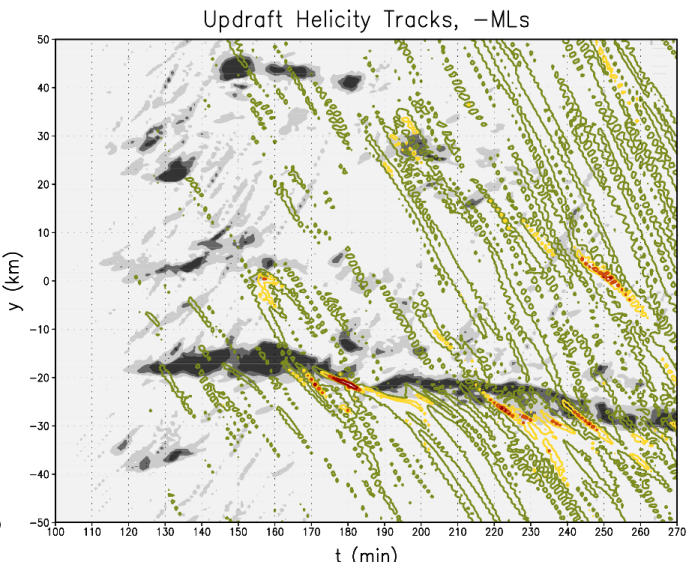
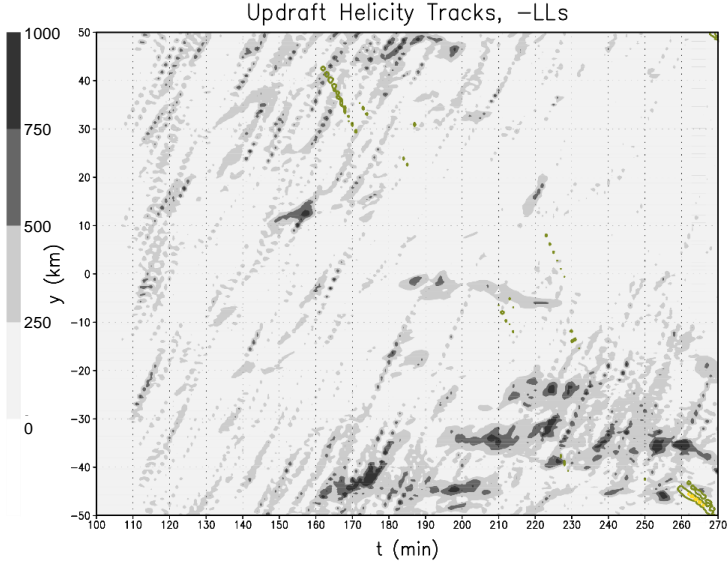
Idealized simulations: Results

Black contours: 1-km vertical velocity
White contours: 10-m vertical vorticity
Reflectivity (dBZ) | Surface θ' (K)

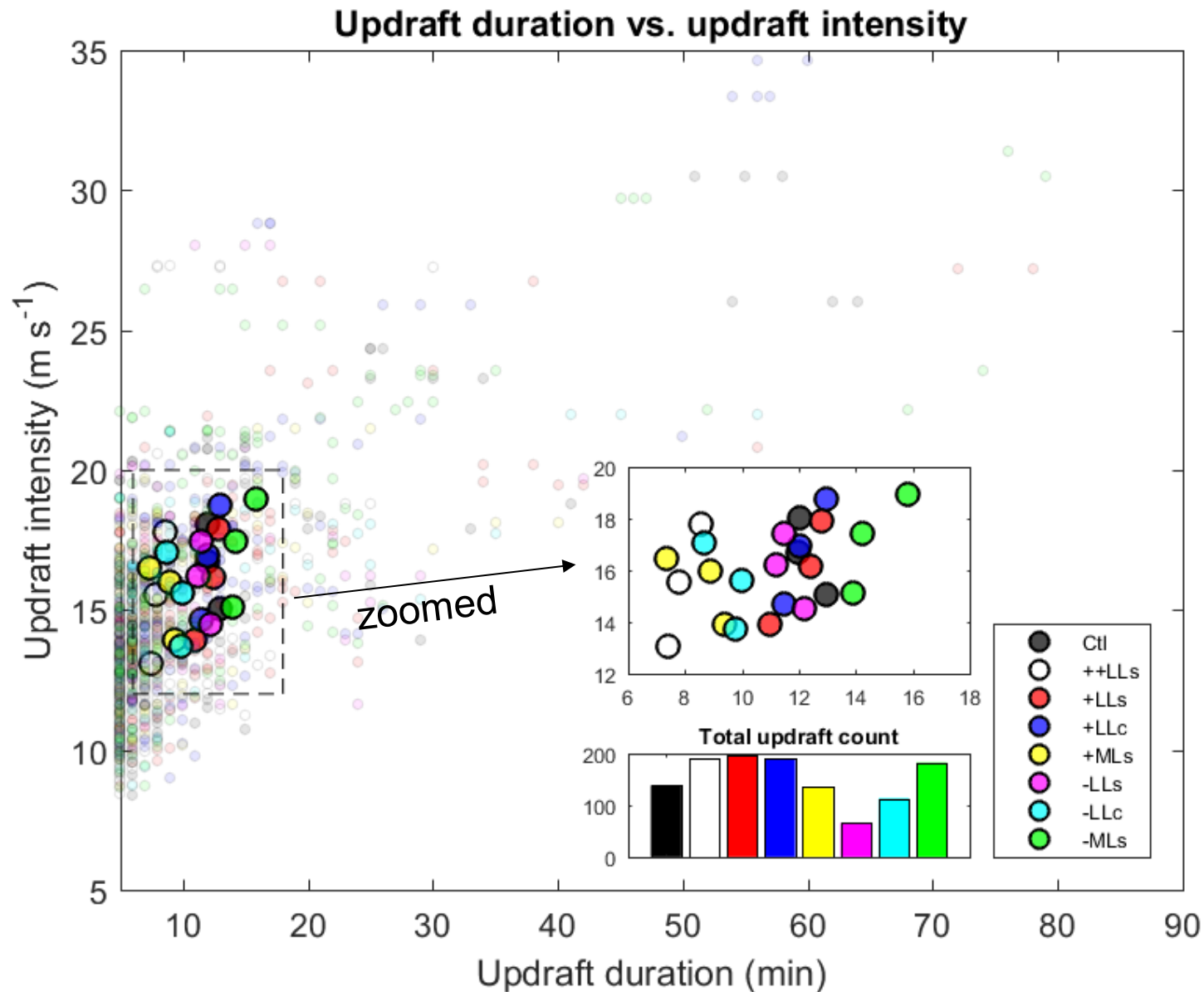


Idealized simulations: Results

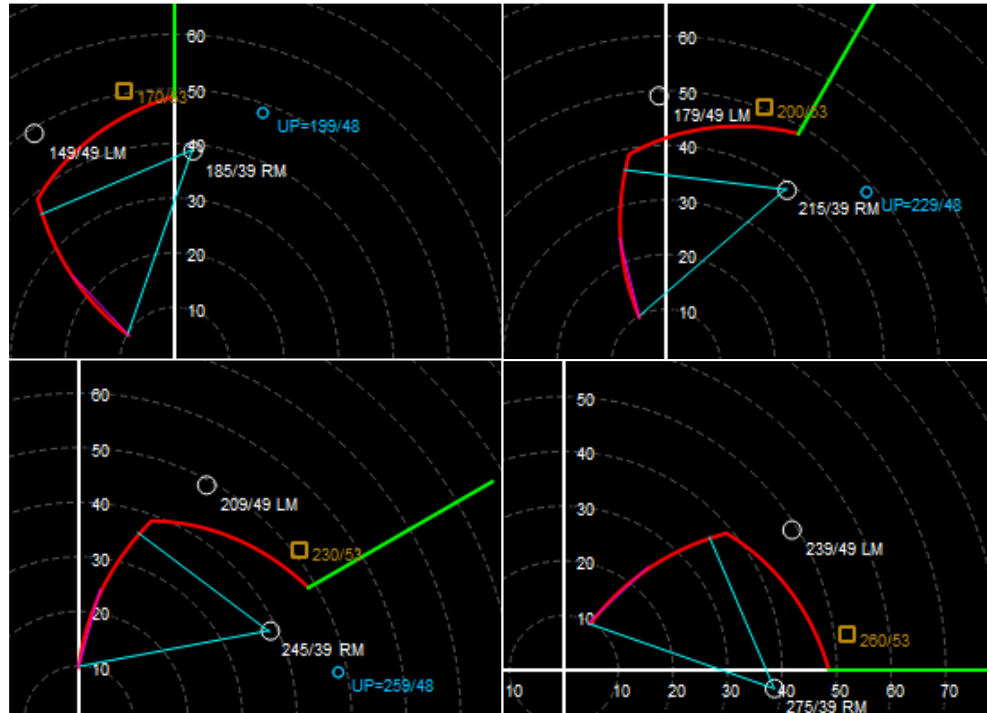
$\zeta = 0.02 \text{ s}^{-1}$ 0.04 s^{-1} 0.06 s^{-1} 0.08 s^{-1}



Idealized simulations: Results

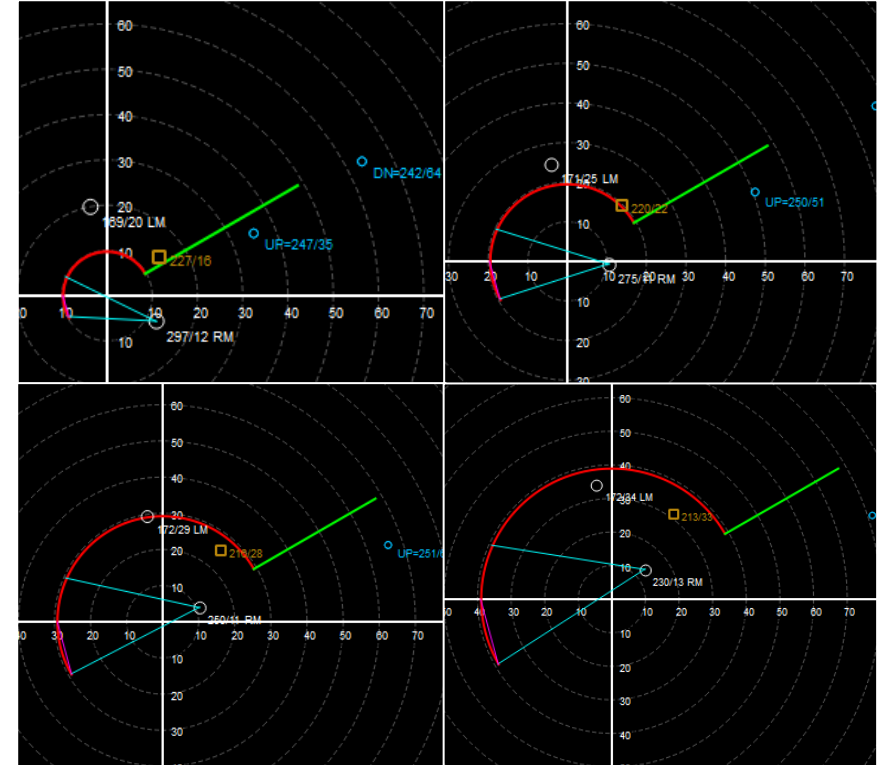


Idealized simulations: Ongoing tests



Maintain hodograph shape, vary orientation relative to initiating boundary

Similar to prior hodograph sensitivity tests, but based upon control hodograph in this matrix



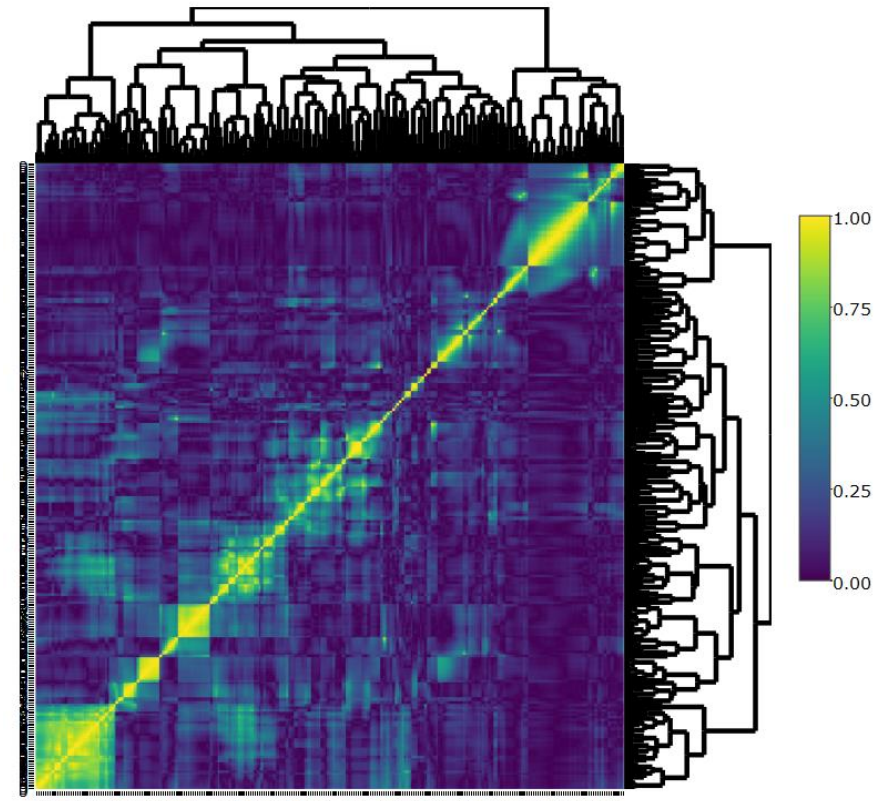
Maintain hodograph shape, vary diameter of "ball cap"

Statistical modeling: Research questions

Two primary questions:

1. Are there statistically significant differences between a tornado-producing environment and an null environment at operational grid lengths, as described by a set of pre-selected variables in HSLC severe environments?
2. What statistical techniques (and in what order/combination) can identify predictive variables, as well as determine the variables' corresponding weights of influence, to differentiate severe/nonsevere environments?

Techniques include: Clustering, linear regression, development of statistical models

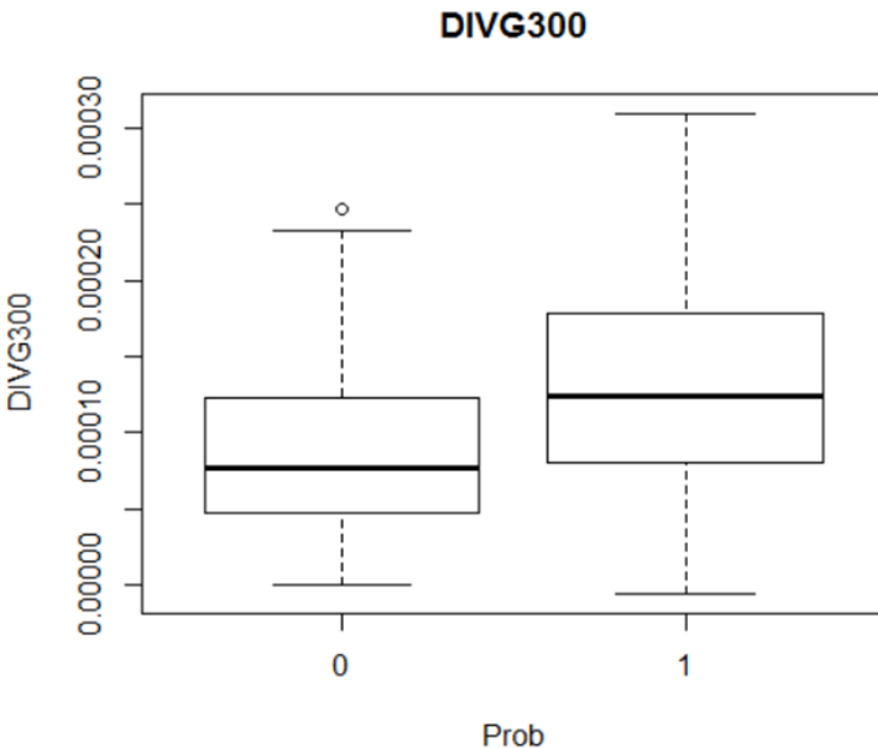


Clustering technique

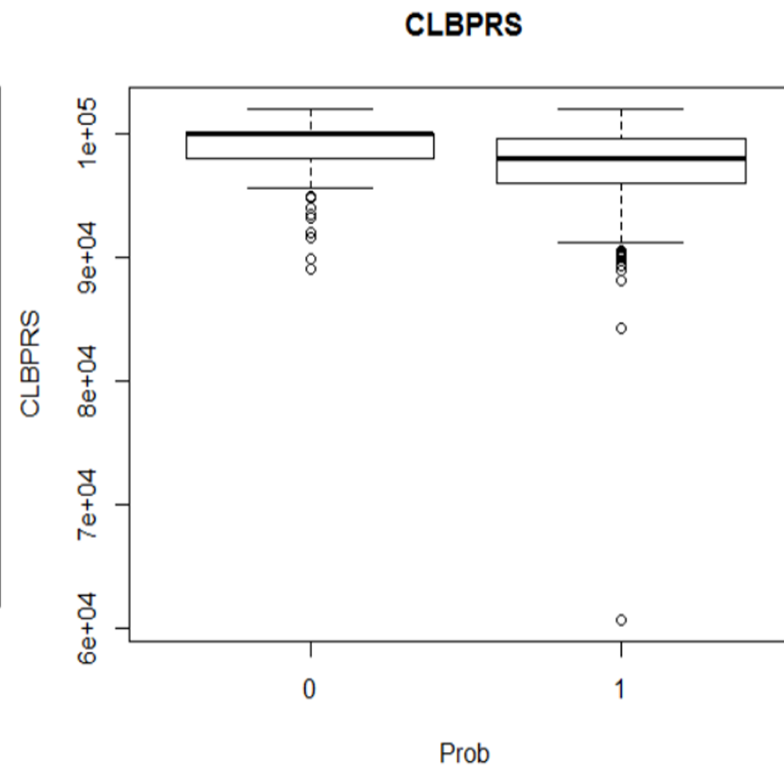
Statistical modeling: Preliminary results

Discriminators that were identified across several techniques/datasets

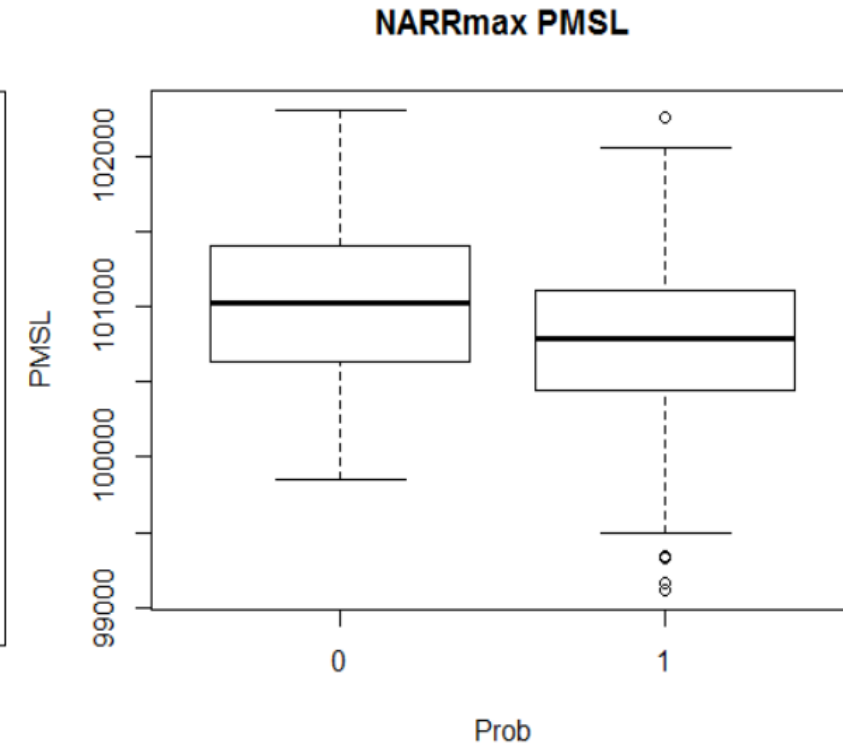
0 = null
1 = tornado report



Proxy for upper-level forcing?



Proxy for low-level stability?



Proxy for synoptic-scale forcing/shear/advection?

Summary

Five components:

- Composite maps and parameters (Sherburn et al. 2017):
 - Favorable environment coupled with strong synoptic-scale forcing for ascent critical in discriminating between severe/nonsevere
 - Combined factors distilled into MOSHE parameter available in beta form on SPC Mesoanalysis; adjustments maybe coming
- Process studies 1: Case simulations to study mesoscale/synoptic scale evolution (King et al. 2017)
 - Rapid destabilization evident in narrow temporal/spatial zone ahead of severe convection
 - Low-level θ_e advection and/or release of potential instability responsible for this destabilization
- NWP studies: Case simulations to investigate resolution requirements
 - 3.6-km, 1.2-km, and 400-m domains all represent convective mode and structure fairly well
 - From 3.6-km to 1.2-km grid spacing, potential convergence of solution
- Process studies 2: Idealized simulations to study convective-scale dynamics
 - Sensitivity studies aimed at understanding prior environmental discriminators
 - Increasing low-level shear or lapse rates appears to increase potential for strong low-level updrafts and vortices to interact
- Statistical studies: Dynamical-statistical downscaling to investigate predictability
 - In progress; several factors indicate large-scale forcing and low-level stability again among most important considerations